

Astrobiology

Lecture 8

Planetary climates and habitable zones

Trieste University, Academic Year 2021-2022
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Climate models and planetary habitability

- The determination of the physical conditions at the planetary surface requires the use of climate models
 - With climate models we can take into account the greenhouse effect and a variety of other processes that affect the surface planetary conditions
- Climate models, originally developed for Earth studies, are becoming a key tool for modeling planetary habitability
 - The state-of-the-art models, called “Global Circulation Models” (GCM), are extremely time consuming
 - Simplified climate models are often used for studies of planetary habitability

Bibliographic material:

Pierrehumbert (2010) Principles of Planetary Climate

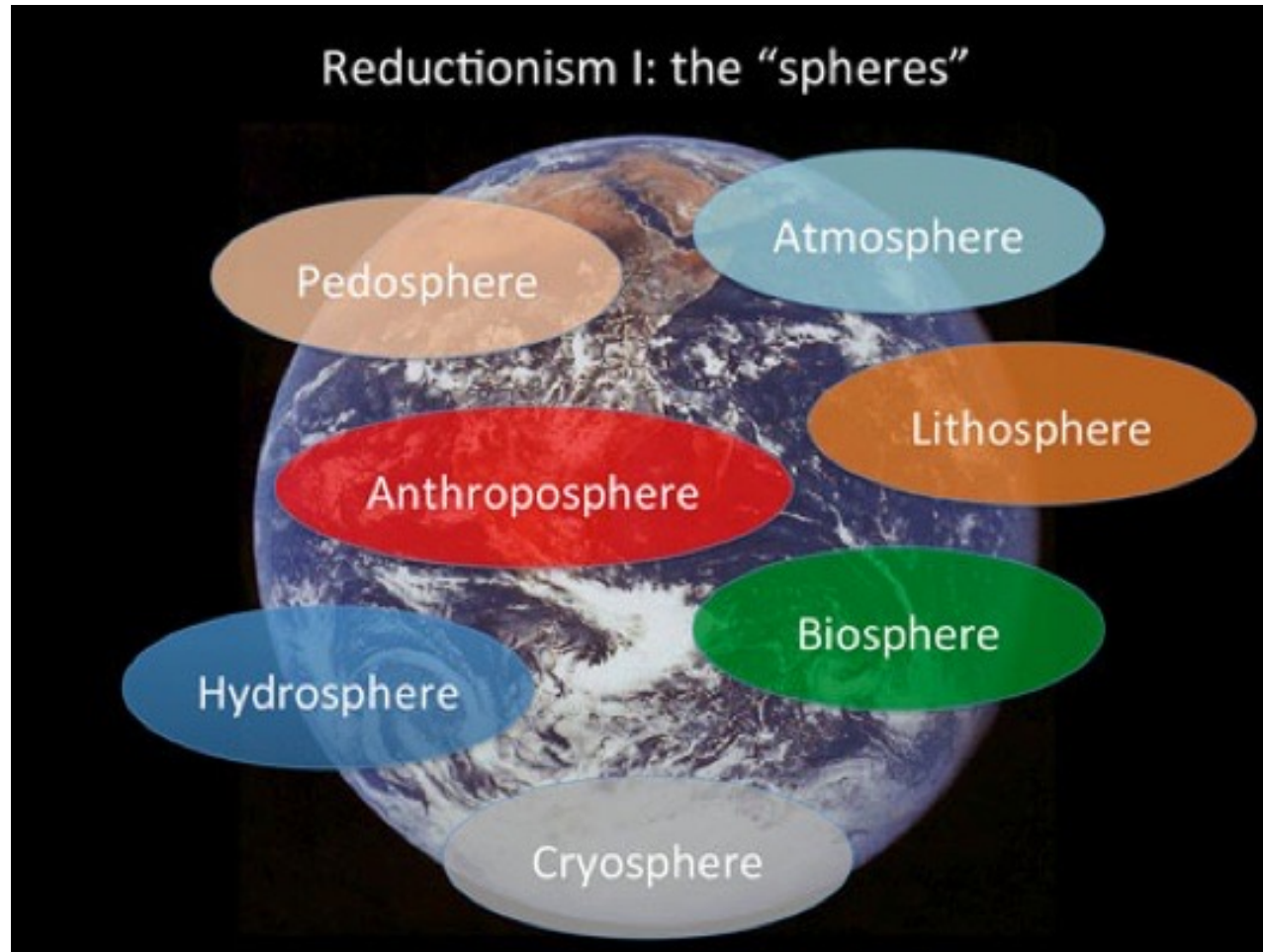
Neelin (2011) Climate Change and Climate Modeling

Complexity of the climate system: the problem of the time scales

Time scales of different components of the climate system

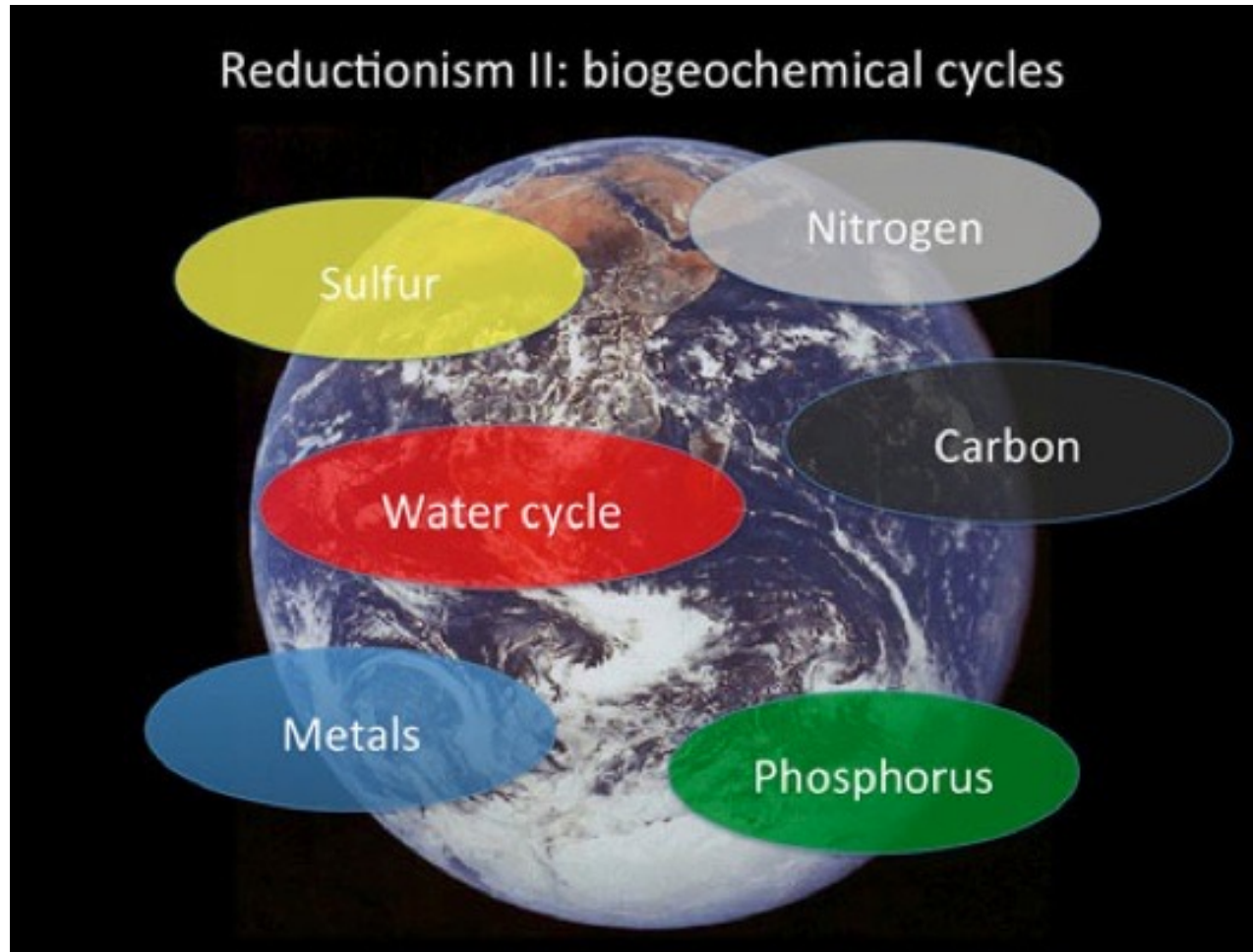
<i>Atmosphere</i>	
Overall response time to heating	months
Typical spin-down time of wind if nothing is forcing it	days
Frontal system lifetime (1000s of km)	days
Convective cloud lifetime (100 m to km horizontal; up to 10 km vertical)	hours
Time scale for typical upper-level wind (20 m s^{-1}) to cross continent (a few 1000 km)	days
<i>Ocean</i>	
Response time of upper ocean (above thermocline) to heating	months to years
Response time of deep ocean to atmospheric changes	decades to millennia
Ocean eddy lifetime (10s to 100 km)	months
Ocean mixing in the surface layer	hours to days
Time for typical ocean current (cm s^{-1}) to cross ocean (1000s of km)	decades
<i>Cryosphere</i>	
Snow cover	months
Sea ice (extent and thickness variations)	months to years
Glaciers	decades to centuries
Ice caps	centuries to millennia
<i>Land surface</i>	
Response time to heating	hours
Response time of vegetation to oppose excess evaporation	hours
Soil moisture response time	days to months
<i>Biosphere</i>	
Ocean plankton response to nutrient changes	weeks
Recovery time from deforestation	years to decades
<i>Lithosphere</i>	
Isostatic rebound of continents (after being depressed by weight of glacier)	10 000s of years
Weathering, mountain building	1 000 000s of years

Reductionistic approaches to tackle the complexity of the climate system



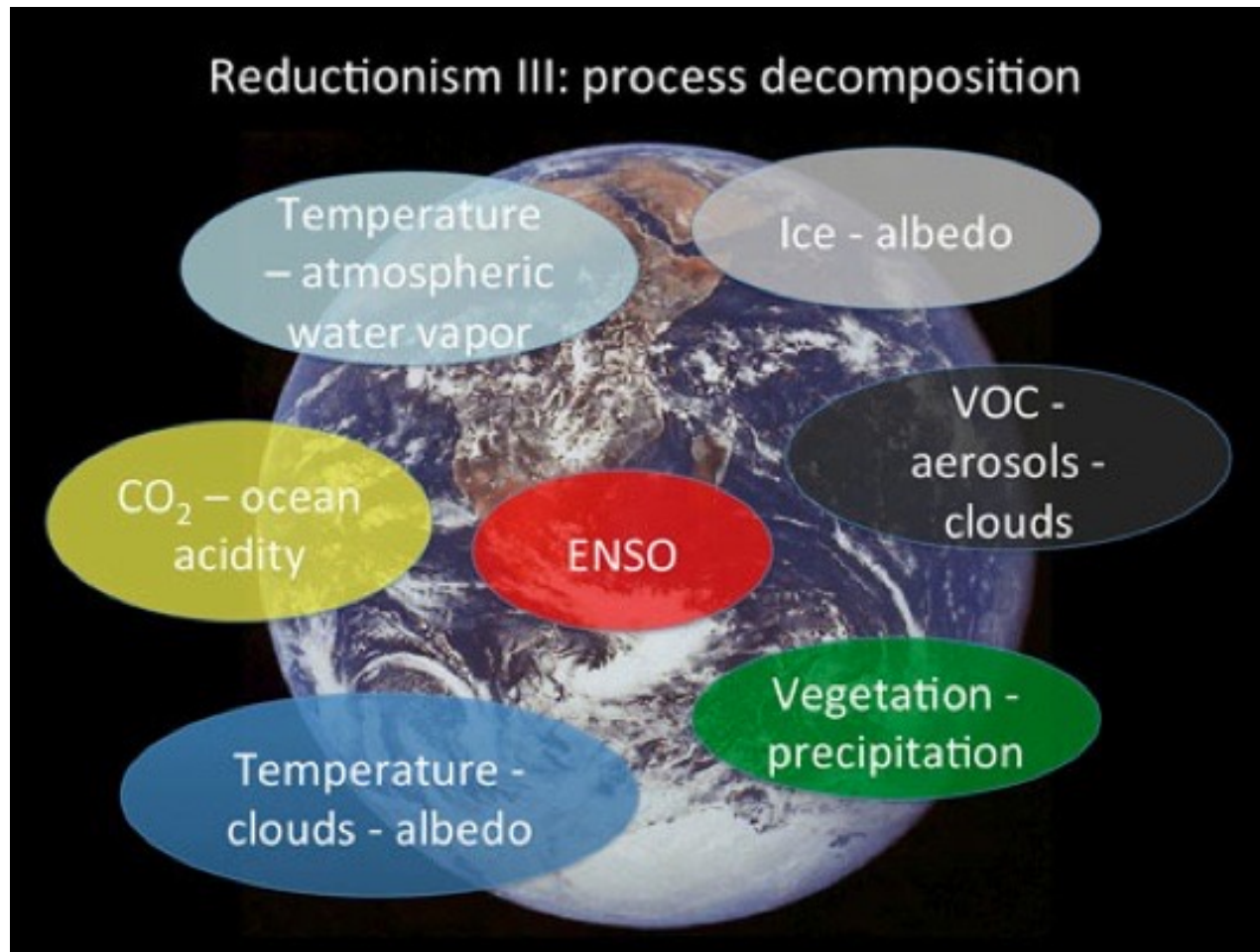
Provenzale (2013)

Reductionistic approaches to tackle the complexity of the climate system



Provenzale (2013)

Reductionistic approaches to tackle the complexity of the climate system

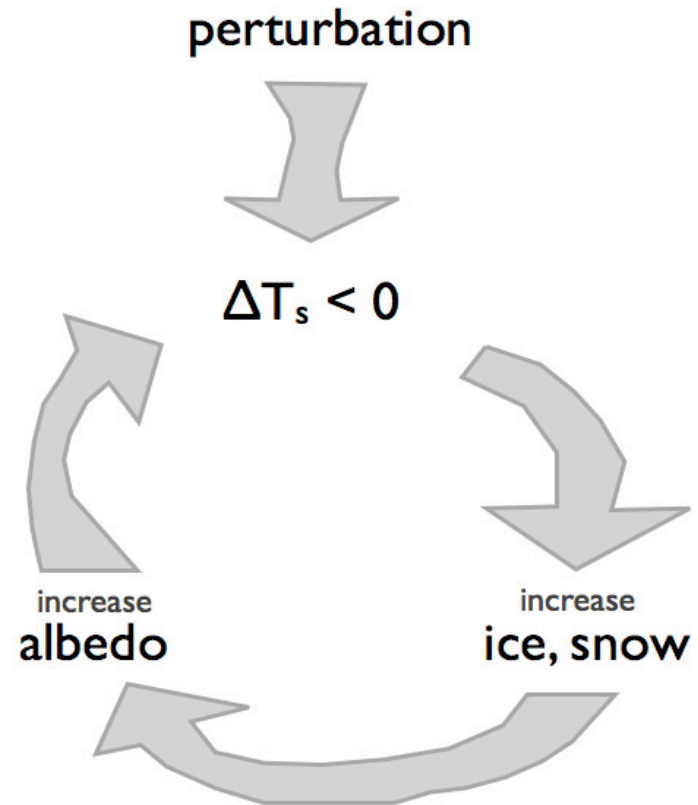


Provenzale (2013)

Climate instabilities

can drive the planet temperature out of the range of habitability

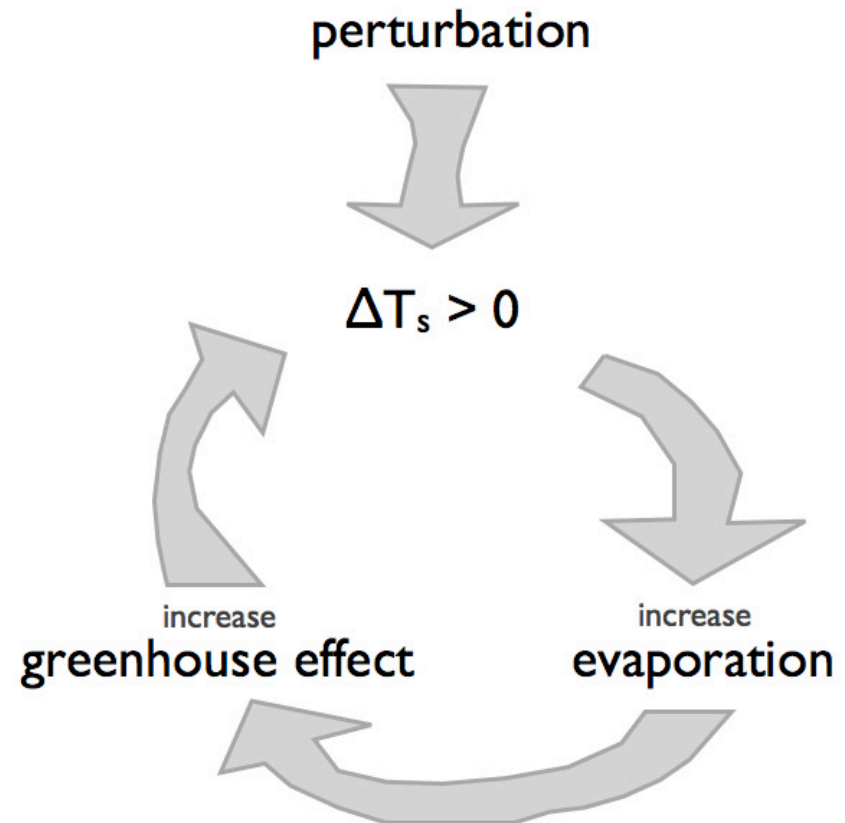
- **The ice-albedo feedback**
 - If water and ice are present on the planet, a decrease of temperature will increase the surface covered by ice
 - Ice has a very high albedo and an increase of the ice extension will cool the planet even more
 - In extreme conditions, this feedback may lead to a “snowball planet”, i.e. a planet fully covered by ice



Climate instabilities

can drive the planet temperature out of the range of habitability

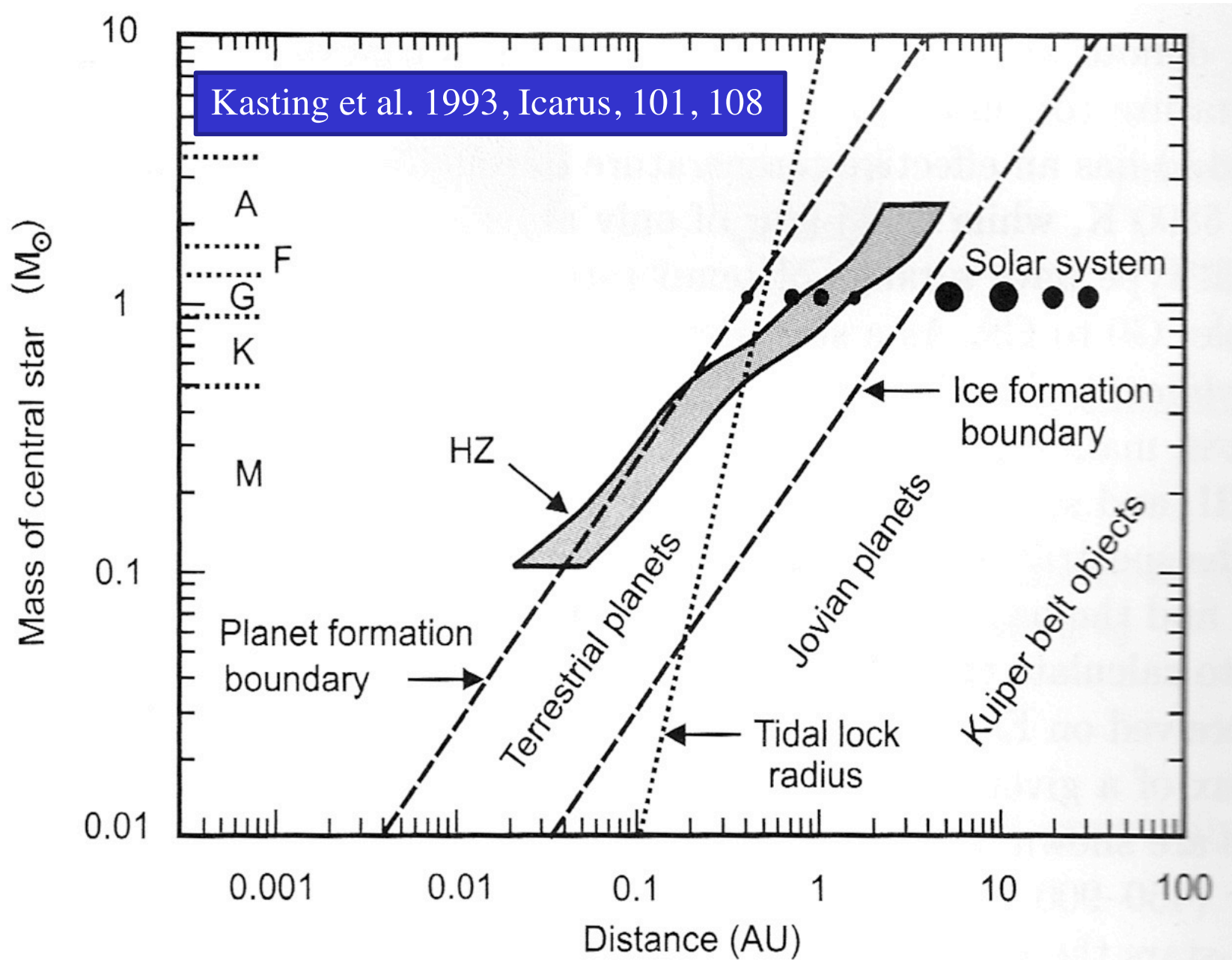
- **The temperature-water vapour feedback**
 - If water is present on the planet, the water vapour pressure rises with temperature
 - Water vapour is a strong greenhouse gas, and a rise of water vapour will rise the temperature even more
 - In normal conditions, this feedback is not catastrophic because the cooling rate scales as σT^4



The classic habitable zone

- Early calculations of planetary habitability were performed before exoplanets were discovered
 - J. Kasting and collaborators (Penn State University)
- Simplified climate models
 - Radiative-convective transport in a single atmospheric column
- Calculated for stars of different spectral types
 - The energy distribution of the stellar spectrum affects the albedo
- Definition of the circumstellar habitable zone
 - Interval of distances from the central star where a planet can have surface conditions suitable for the long-term presence of liquid water
 - Two different criteria are adopted to define the inner and outer edge of the habitable zone

The classic habitable zone



The inner edge of the habitable zone

- **The runaway greenhouse mechanism**
 - If the temperature-water vapour feedback is extreme, the vapour may reach the outer layers of the atmosphere
 - In the outer layers the water molecules can be dissociated by high energy stellar photons
 - The hydrogen produced by photodissociation can be lost to space
 - This chain of events is called the runaway greenhouse mechanism
 - In the long term, this mechanism may lead to the disappearance of liquid water on the planet
 - The “runaway greenhouse” mechanism is used to define the inner edge of the habitable zone

The outer edge of the habitable zone

- An increase of greenhouse gases in the planetary atmosphere makes the planet habitable at lower levels of stellar flux, i.e. at larger distances from the central star
- To keep the planet habitable in the outer regions of the habitable zone it is assumed that the planetary atmosphere is dominated by CO₂
 - as in the case of Mars
- The amount of CO₂ that is able to warm the planet at low levels of insolation is limited by the onset of CO₂ clouds with high albedo, which would counteract the heating due to greenhouse effect
- The outer edge is defined via the “maximum greenhouse” criterion, i.e. the maximum amount of CO₂ before cooling of the clouds take place

Mechanisms of climate stabilization

In the definition of the classic habitable zone it is assumed that the planet has the capability of adjusting its level of CO₂ through a mechanism of climate stabilization

The fact that Earth's climate has been relatively stable in the course of geological time scales suggests the existence of a mechanism of climate stabilization

This mechanism must have been able to stabilize the Earth's climate despite changes that have occurred in terms of solar radiation, atmospheric composition and other factors

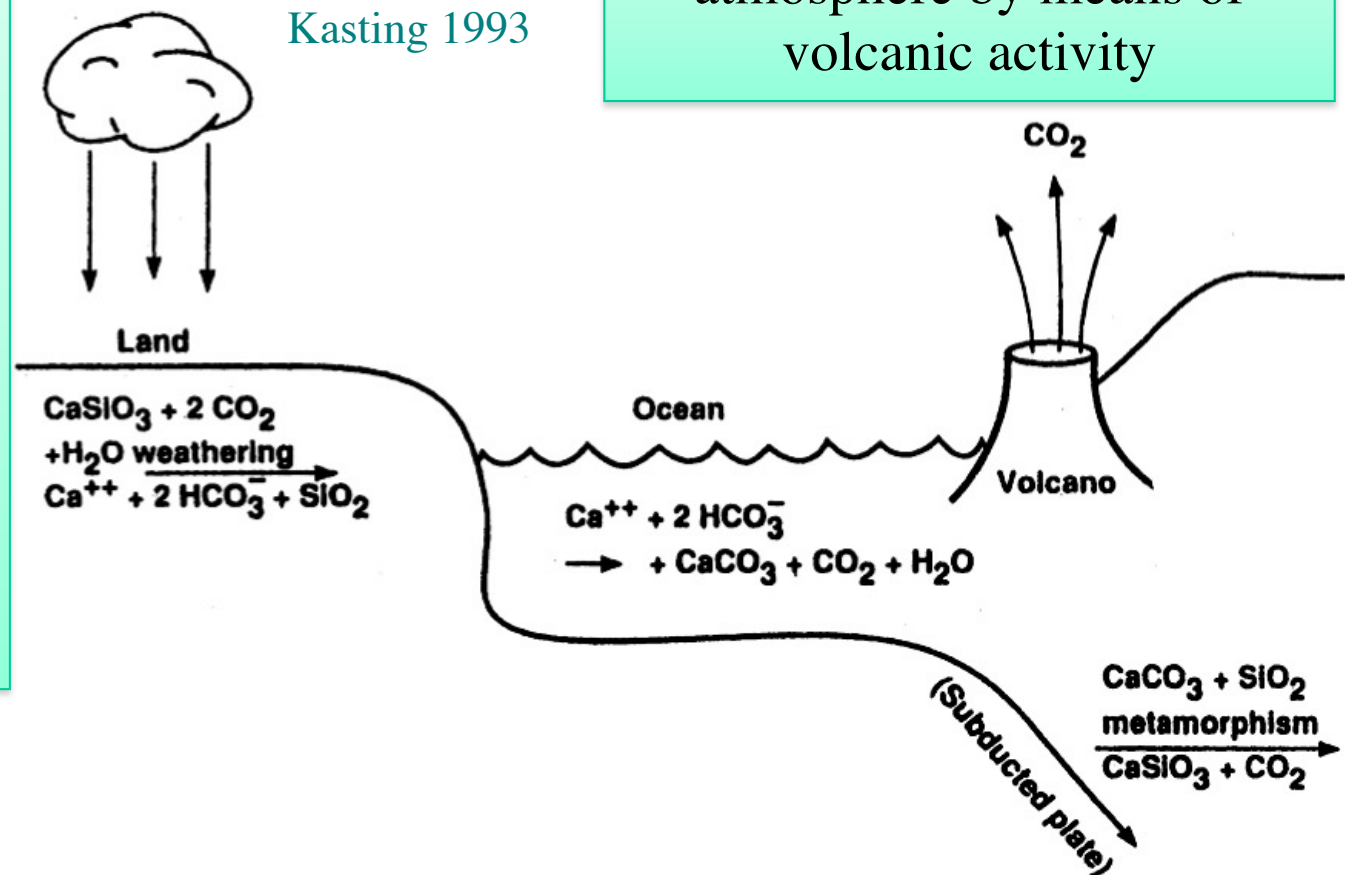
The mechanism invoked for the Earth is based on a CO₂ inorganic cycle

The CO₂ cycle of climate stabilization

1. - Weathering processes remove CO₂ from the atmosphere

2. - The chemical products are gradually deposited to the bottom of the oceans and eventually subducted, due to tectonic activity

3. - CO₂ from the Earth's mantle is emitted to the atmosphere by means of volcanic activity



The CO₂ cycle of climate stabilization

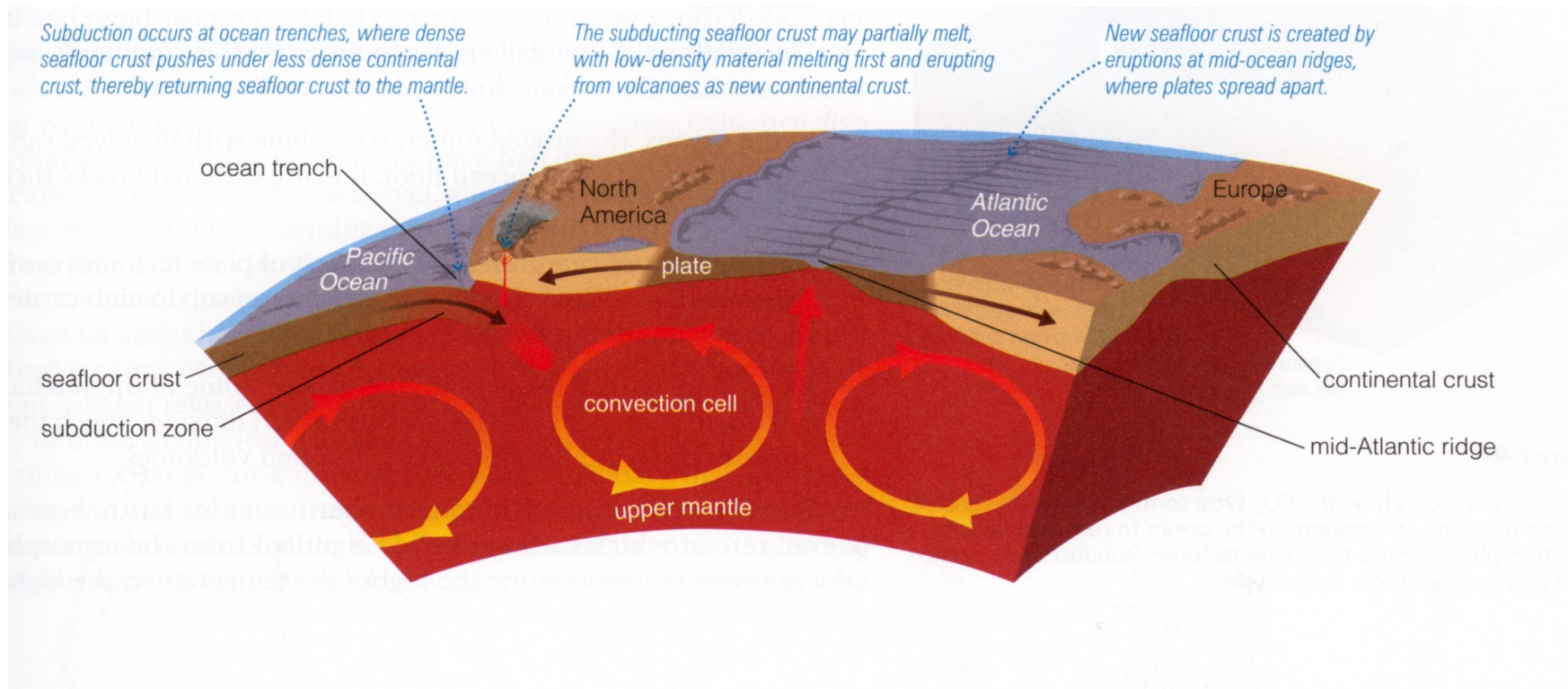
The weathering efficiency increases with atmospheric temperature

The rate of CO₂ emission is independent of the atmospheric temperature

As a result, there is a negative feedback temperature-CO₂ that stabilizes the climate

- The time scale of the cycle is estimated to be $\sim 5 \times 10^5$ years

- The existence of tectonics and volcanism is necessary for the existence of the CO₂ stabilization cycle
 - In the present-day Solar System, only the Earth has these types of geophysical activities
 - The CO₂ cycle of climate stabilization is invoked in the definition of the classic habitable zone
 - How common this mechanism can be in terrestrial-type exoplanets is currently the subject of studies of exoplanetary interiors

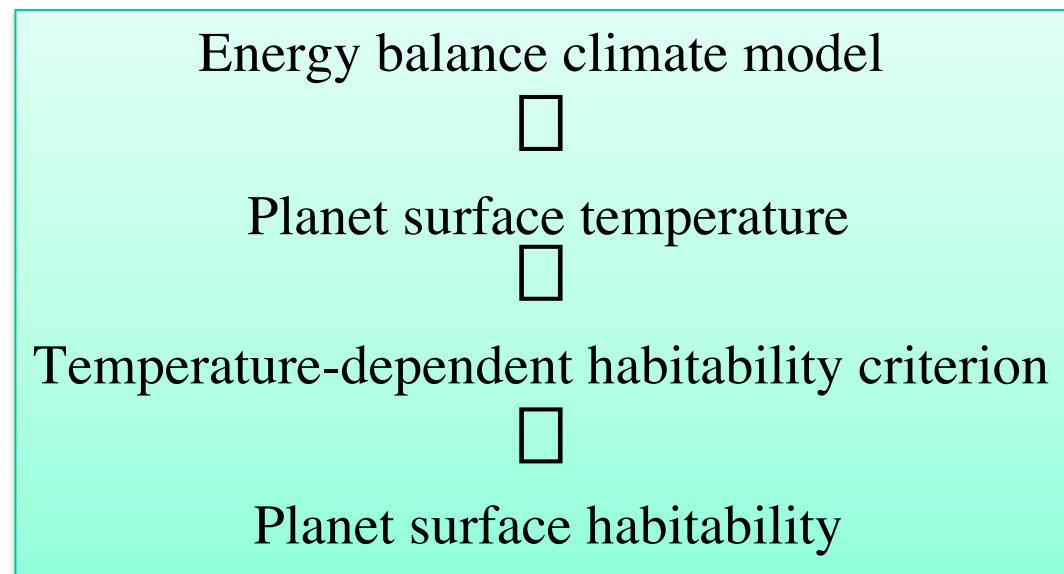


Habitability under the planet surface

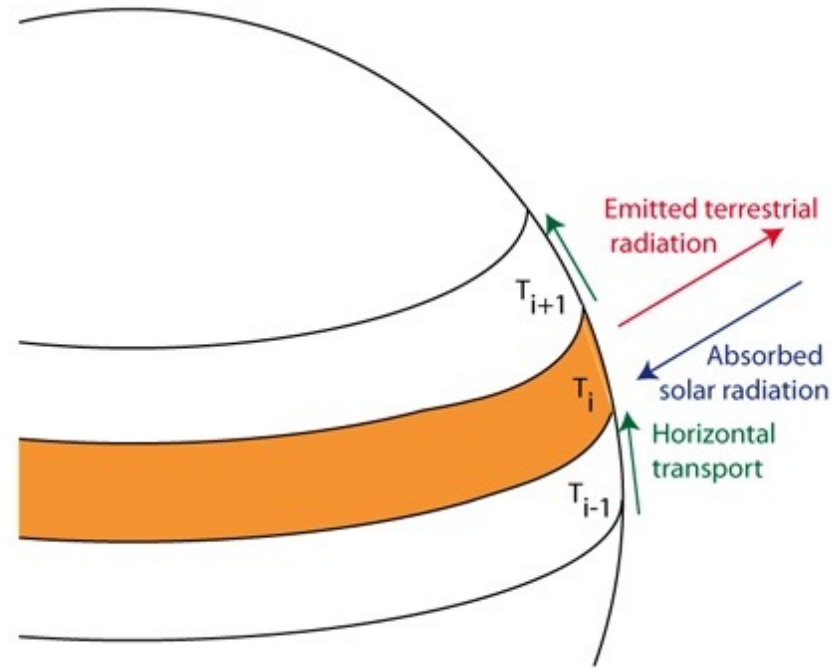
- The definition of habitable zone relies on the concept of surface habitability
 - Habitability under the planet surface could be present in planetary bodies outside the circumstellar habitable zone, in particular beyond the outer edge
- Temperature and pressure gradients may yield conditions of habitability in the interior of planets or satellites
 - Internal sources of heat yield a temperature gradient in the planet interior
 - The pressure gradient towards the planetary interior may improve the conditions of habitability, e.g. by shifting the pressure above the triple point of water

Surface habitability with Energy Balance Models (EBMs)

Simplified climate models aimed at predicting
the seasonal and latitudinal distribution
of the surface temperature



Energy balance models (EBM) of planetary climate



$$I_i + C_i \frac{\partial T}{\partial t} - \frac{\partial}{\partial x} \left[D_i (1 - x^2) \frac{\partial T}{\partial x} \right] = S_i (1 - A_i)$$

$x = \sin \phi$ (ϕ is the latitude)

Brief summary of recent studies of planetary habitability carried out at INAF-OATs

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THE HABITABLE ZONE OF EARTH-LIKE PLANETS WITH DIFFERENT LEVELS OF ATMOSPHERIC PRESSURE

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Earth-like Surface
Temperature model
(ESTM) with
improved treatment
of the latitudinal
transport

MODELING THE SURFACE TEMPERATURE OF EARTH-LIKE PLANETS

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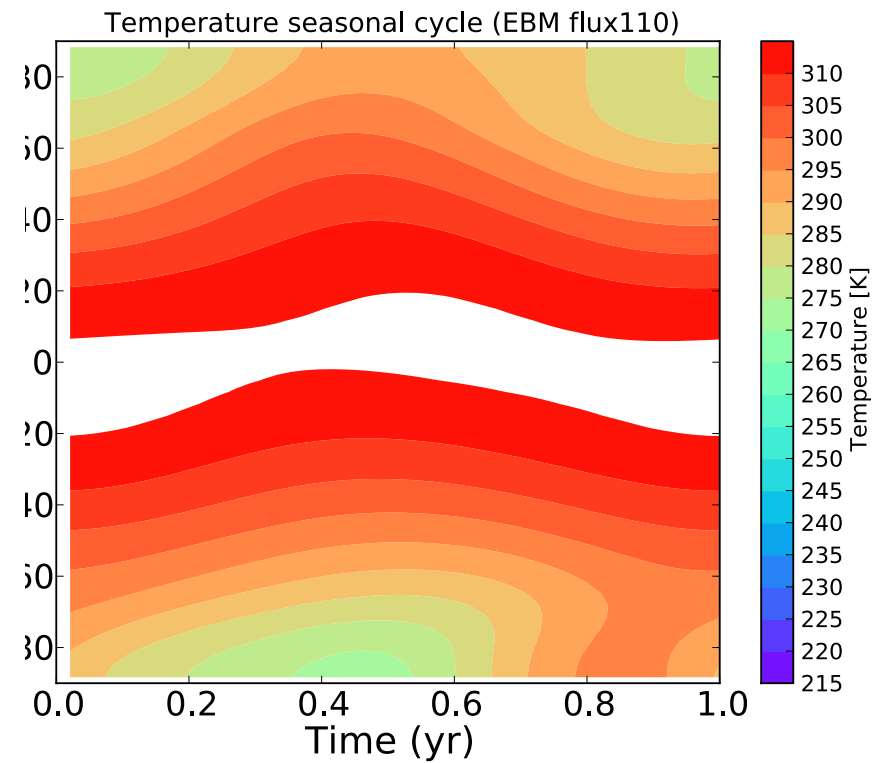
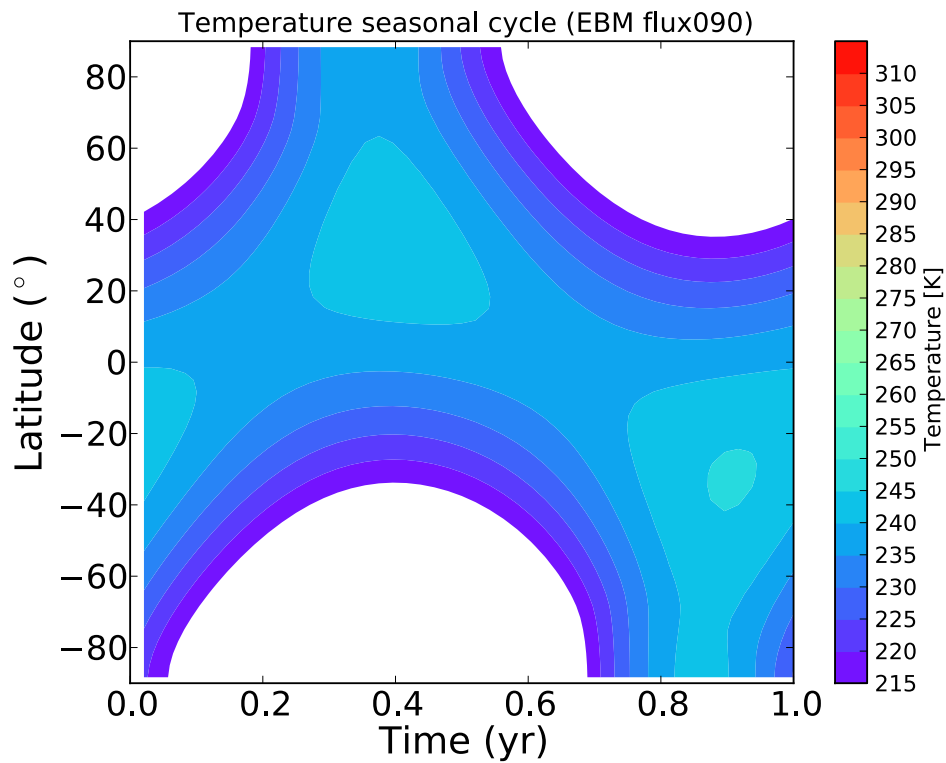
From climate models to planetary habitability: temperature constraints for complex life

Laura Silva¹, Giovanni Vladilo¹, Patricia M. Schulte²,
Giuseppe Murante¹, Antonello Provenzale³

Examples of application of the ESTM

Seasonal and latitudinal surface temperature of the Earth

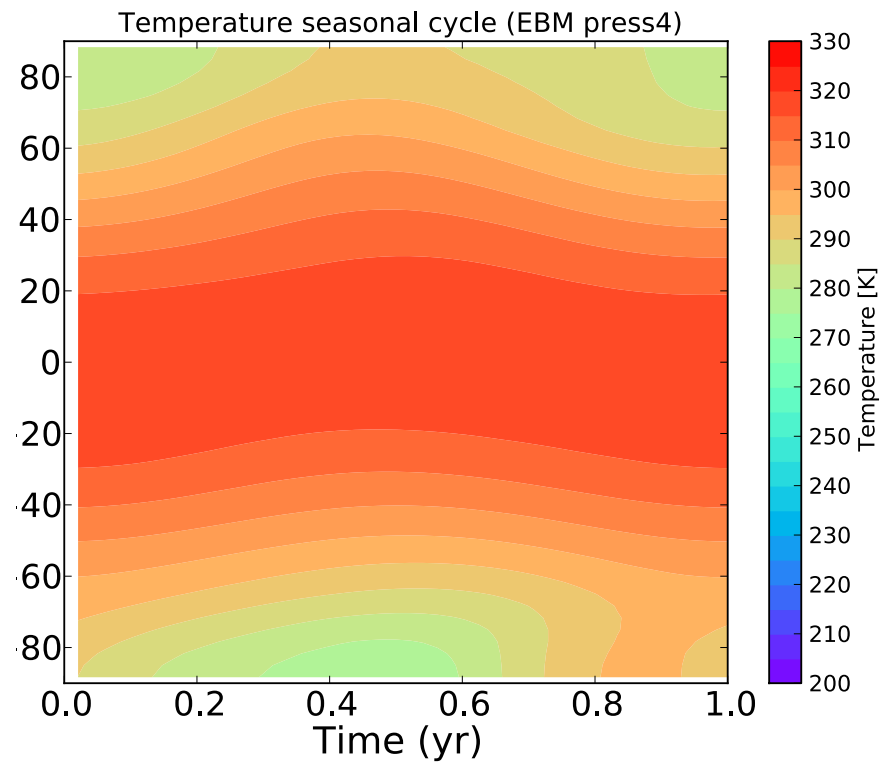
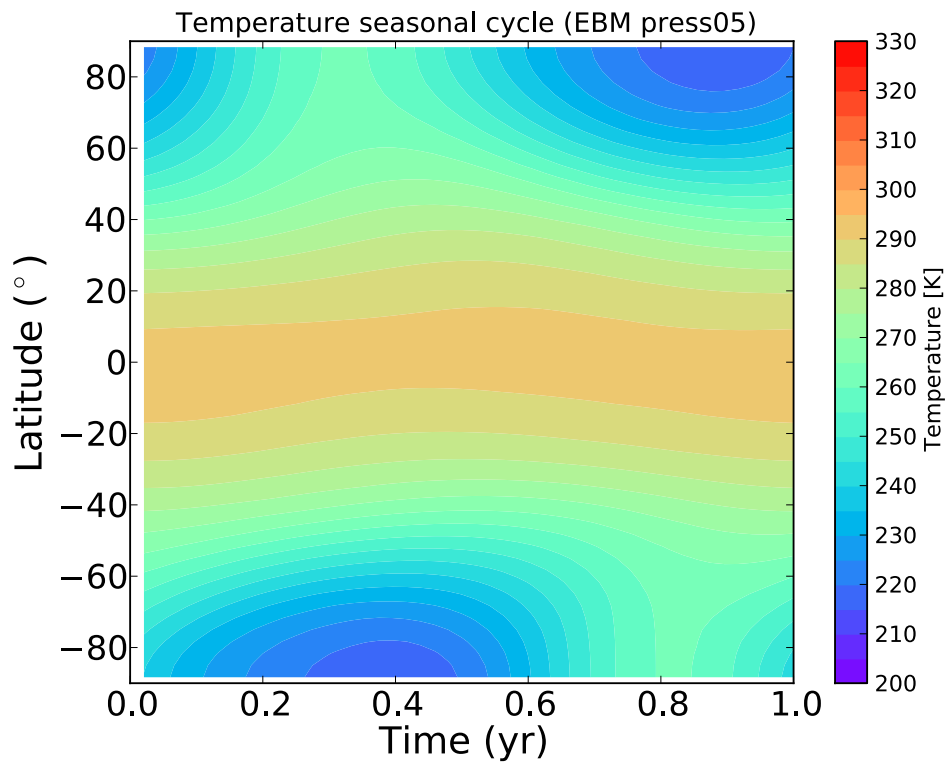
Variation of stellar insolation



Examples of application of the ESTM

Seasonal and latitudinal surface temperature of the Earth

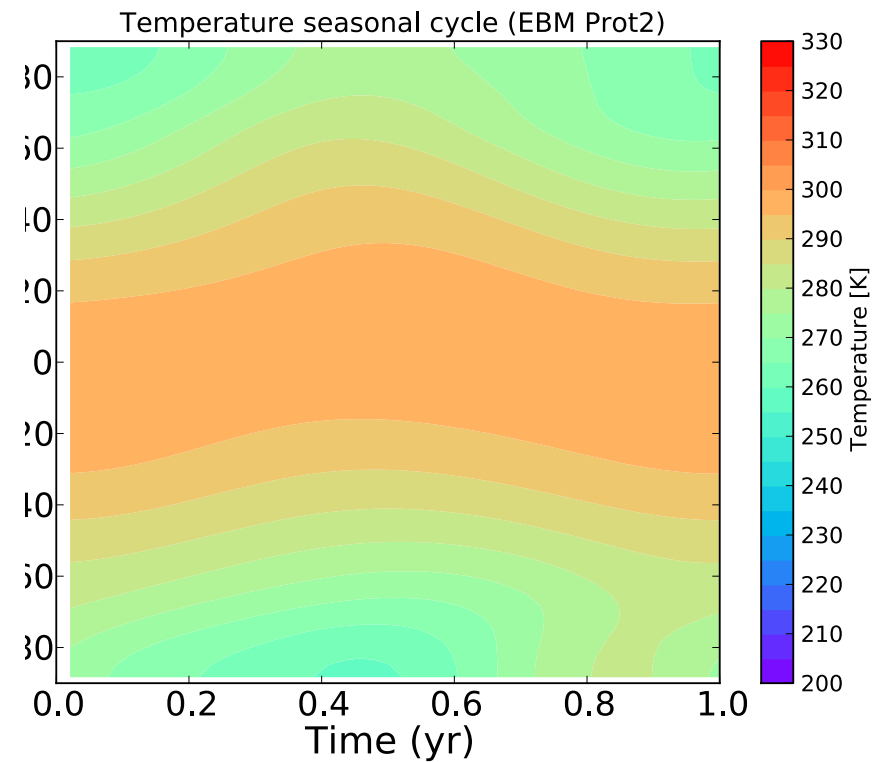
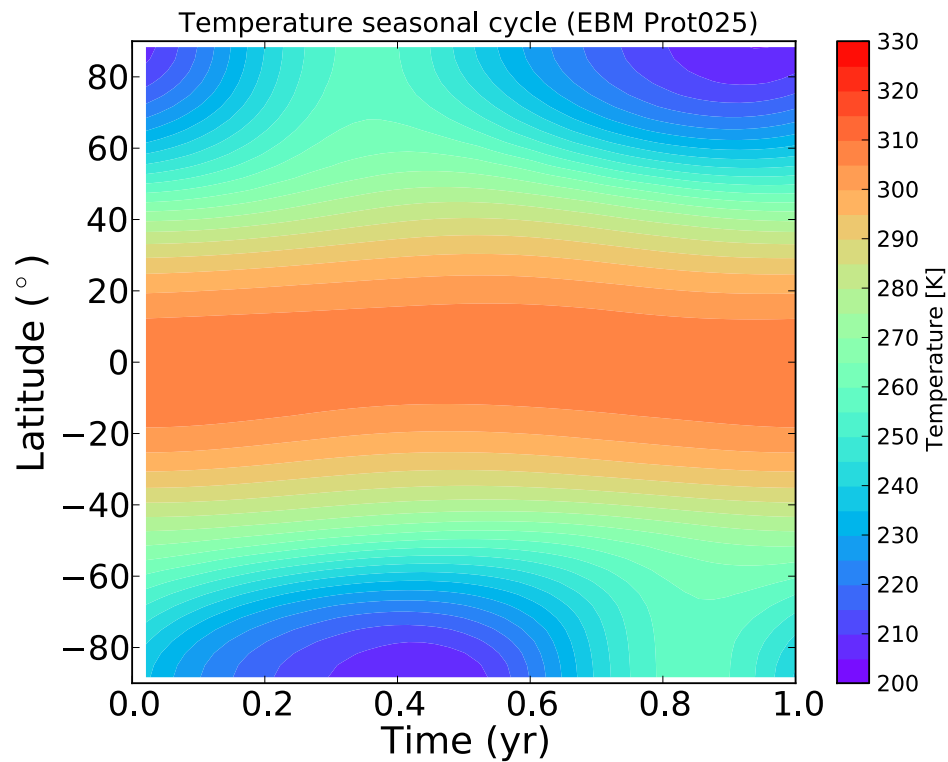
Variation of surface pressure



Examples of application of the ESTM

Seasonal and latitudinal surface temperature of an Earth-like planet

Variation of rotation period



Quantifying the habitability with the liquid water criterion

The climate simulation yields the surface temperature as a function of latitude and time

$$T(\phi, t)$$

Habitability function

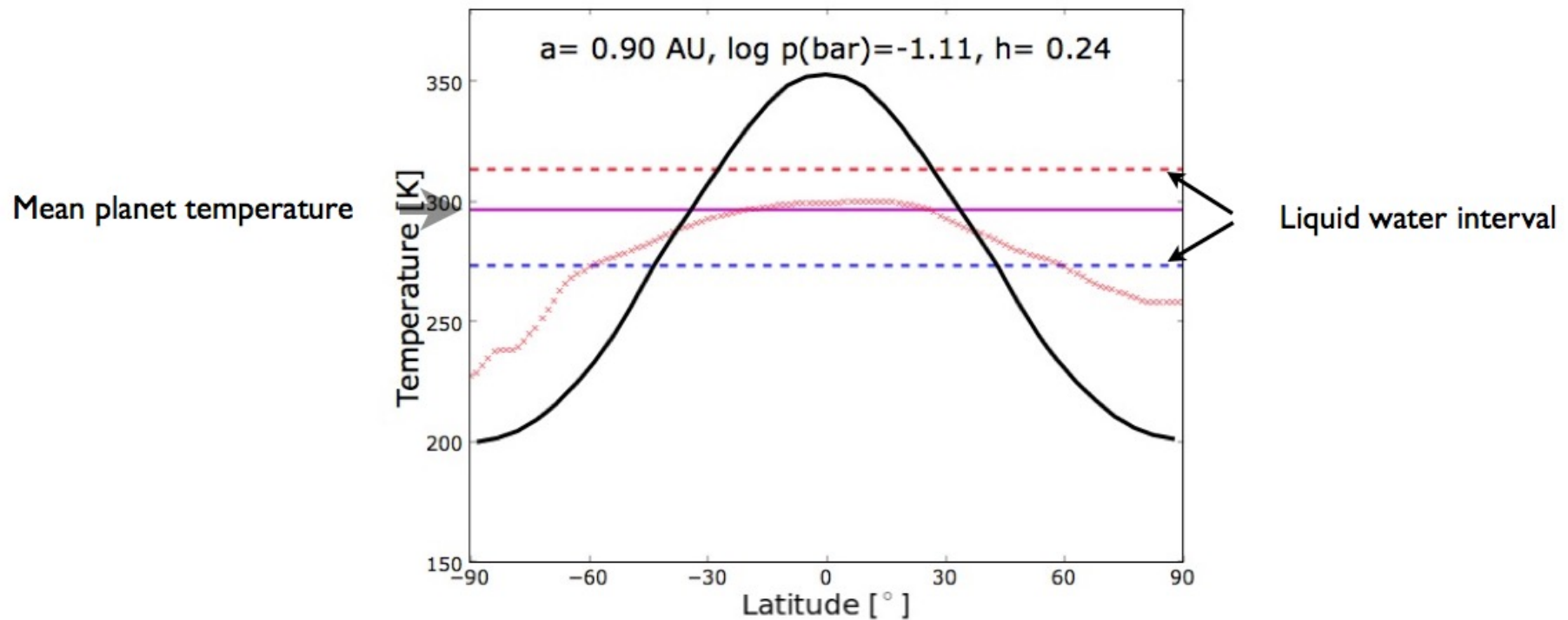
$$H(\phi, t) = \begin{cases} 1 & \text{if } T_{\text{melt}}(p) \leq T(\phi, t) \leq T_{\text{boil}}(p) \\ 0 & \text{otherwise} \end{cases}$$

Liquid water criterion

Mean global annual habitability

$$h = \frac{\int_{-\frac{\pi}{2}}^{+\frac{\pi}{2}} d\phi \int_0^P dt [H(\phi, t) \cos \phi]}{2P}$$

At low pressure,
the mean planet temperature is not a good diagnostic of habitability

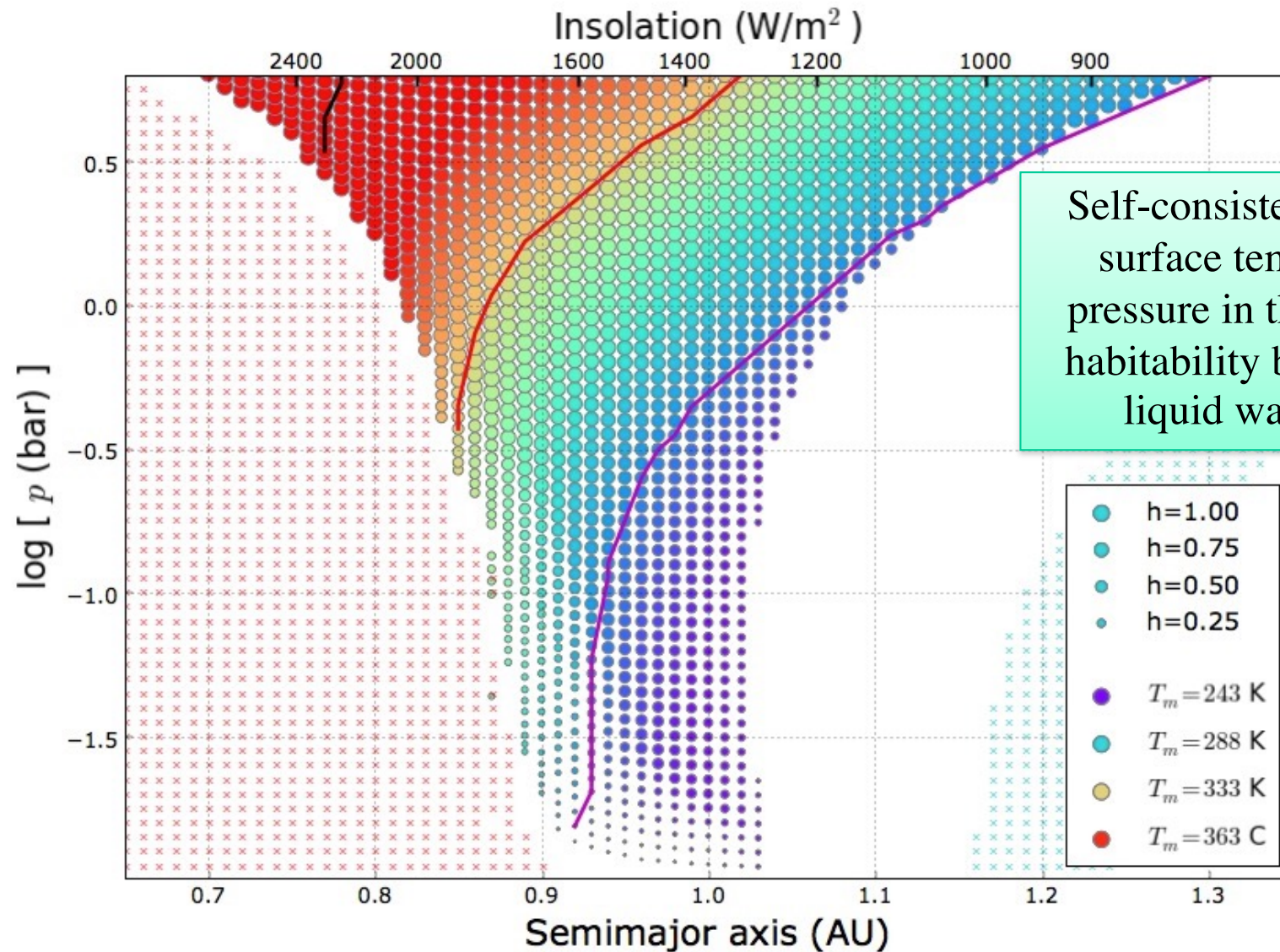


Evaporating planets inside the “habitable zone”

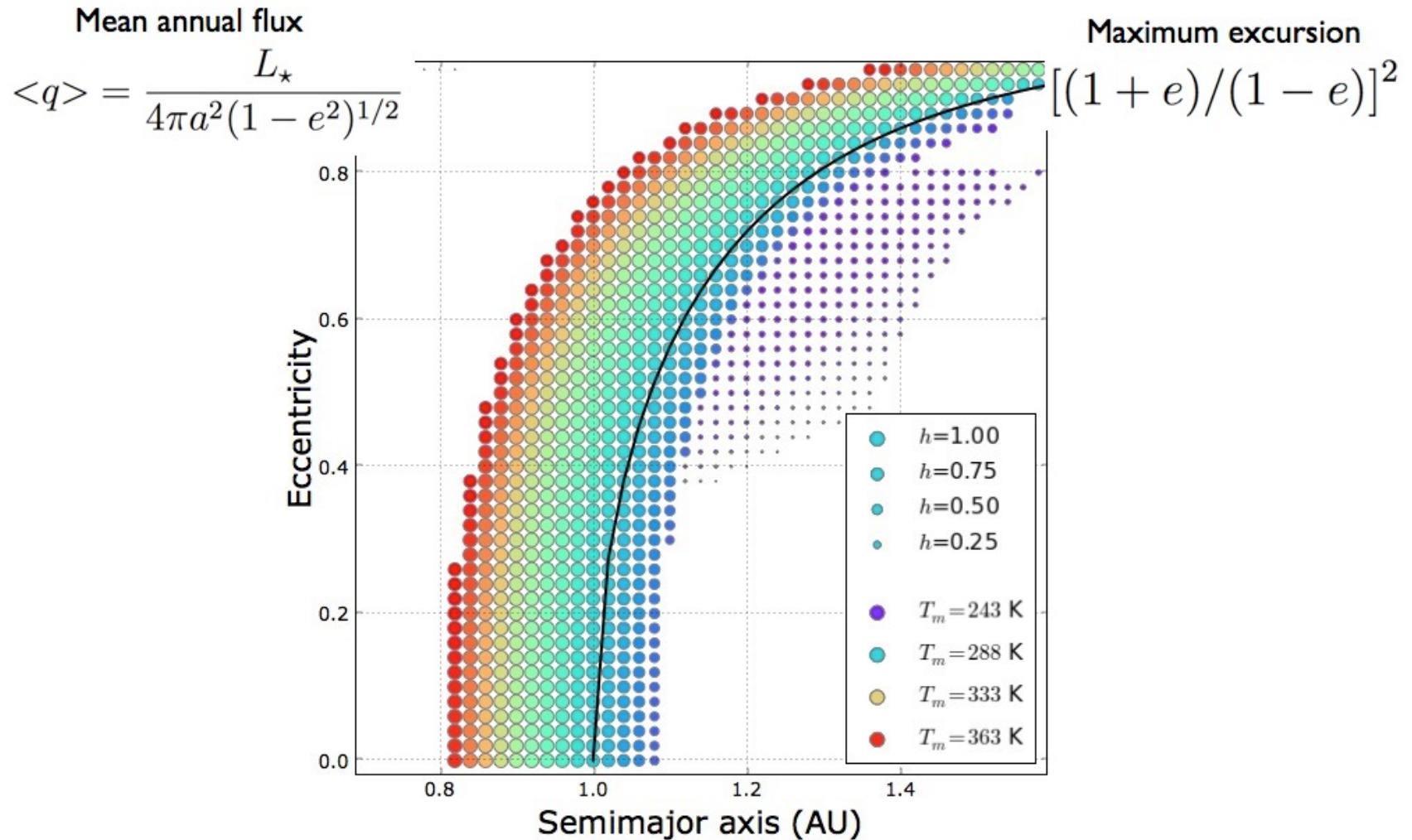
Pressure dependence of the habitable zone

Obtained running a large number of climate simulations

Vladilo et al. (2013)



Effects of eccentricity



The atmospheric mass habitable zone
for complex life and atmospheric biomarkers
Obtained using the temperature limits $0^{\circ}\text{C} \lesssim T \lesssim 50^{\circ}\text{C}$
Silva et al. (2017)

