

V Congresso Società Italiana di Astrobiologia  
15-17 September 2015  
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# Disequilibrium in planetary atmospheres and the search for habitability

Eugenio Simoncini

*Astrophysical Observatory of Arcetri - INAF, Firenze, Italy*

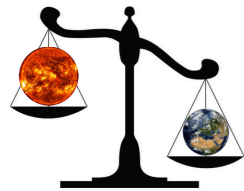
*Co-Chair of the TDE NAI Focus Group*



N A S A  
ASTROBIOLOGY  
I N S T I T U T E

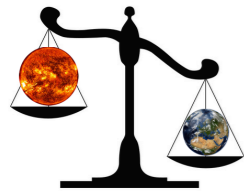
Thermodynamics,  
Disequilibrium and  
Evolution (TDE) – NAI  
Focus Group





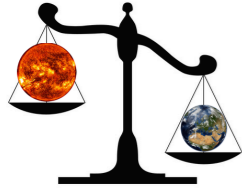
# *Overview of this talk*

- What is chemical disequilibrium
- How to calculate disequilibrium in chemical processes
- Atmospheric Disequilibrium of Earth
- Further applications



## Introduction

*What is chemical disequilibrium, and why should we use it*



[www.nature.com/scientificreports](http://www.nature.com/scientificreports)

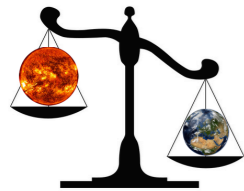
# SCIENTIFIC REPORTS

**OPEN** **Titania may produce abiotic oxygen atmospheres on habitable exoplanets**

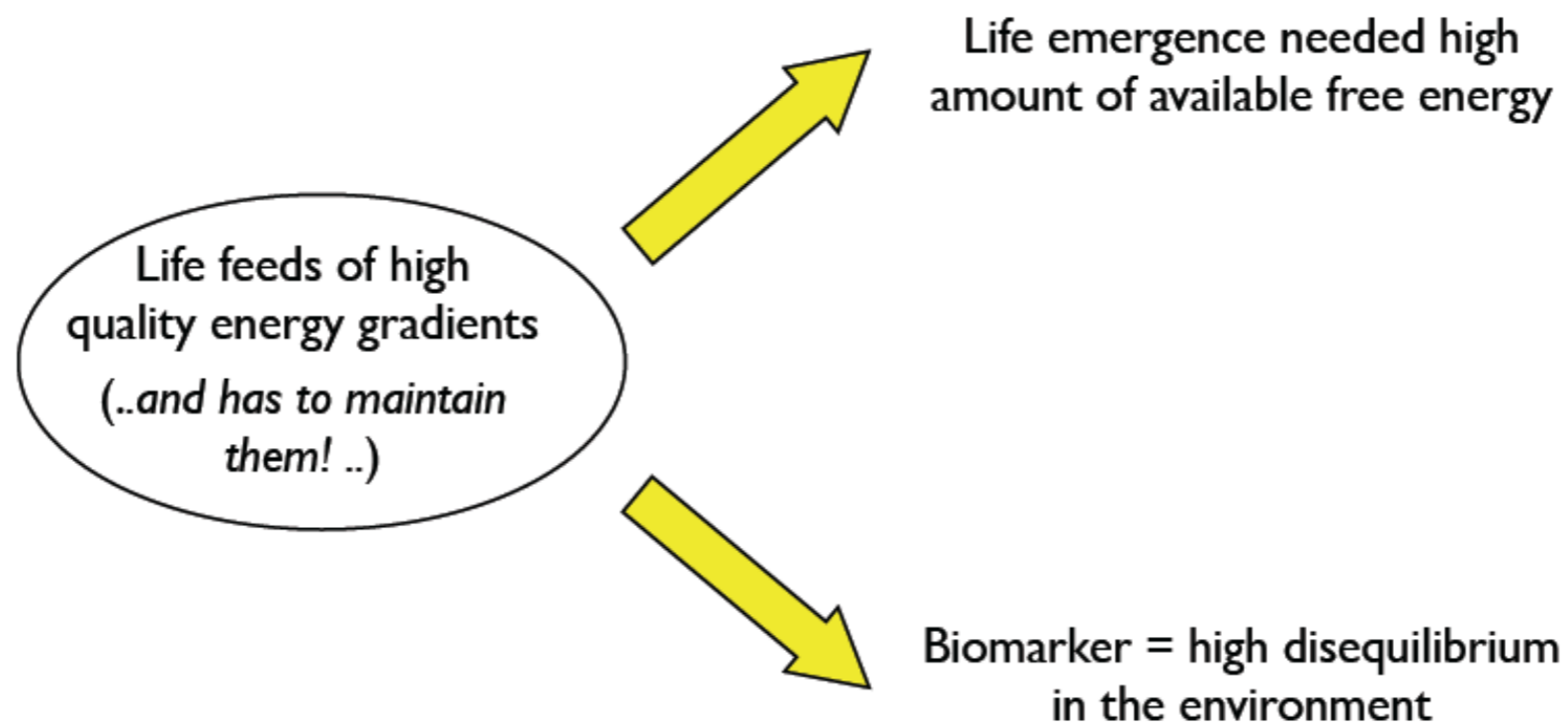
Norio Narita<sup>1,2,3</sup>, Takafumi Enomoto<sup>3,4</sup>, Shigeyuki Masaoka<sup>3,4</sup> & Nobuhiko Kusakabe<sup>2</sup>

Received: 31 March 2015  
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Published: 10 September 2015

The search for habitable exoplanets in the Universe is actively ongoing in the field of astronomy. The biggest future milestone is to determine whether life exists on such habitable exoplanets. In that context, oxygen in the atmosphere has been considered strong evidence for the presence of photosynthetic organisms. In this paper, we show that a previously unconsidered photochemical mechanism by titanium (IV) oxide (titania) can produce abiotic oxygen from liquid water under near ultraviolet (NUV) lights on the surface of exoplanets. Titania works as a photocatalyst to dissociate liquid water in this process. This mechanism offers a different source of a possibility of abiotic oxygen

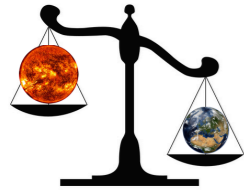


# Life & Disequilibrium



*“The general struggle for existence of animate being is struggle for entropy, which becomes available through the transition of energy from the hot sun to the cold earth” (Boltzmann, 1886)*

*“Life feeds of high quality energy gradient” (Schrödinger, 1944)*



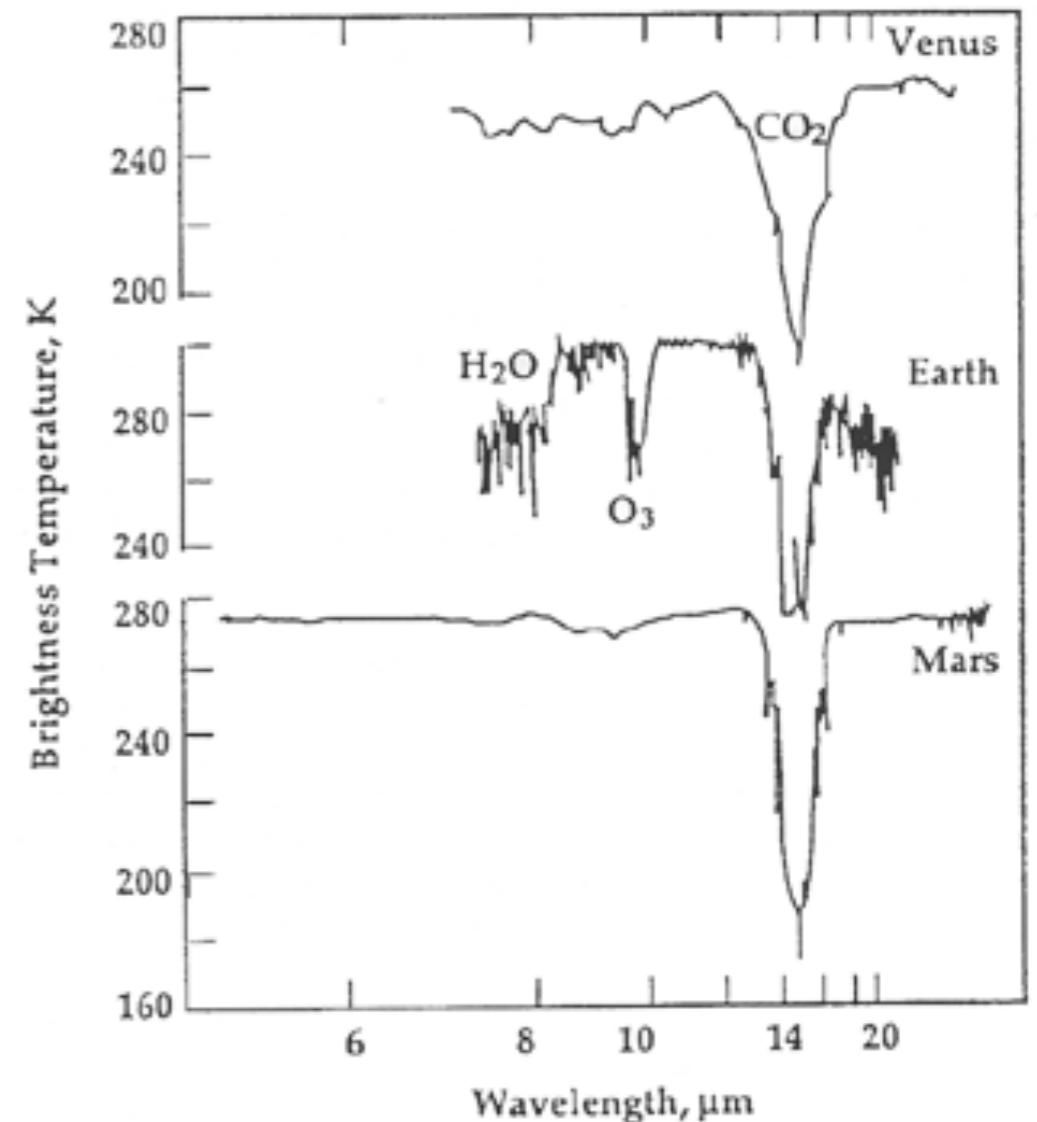
# Life and Earth disequilibria



*The emergence of life allowed the use of more degrees of freedom associated to geological and atmospheric cycles, and consequently the generation of more free energy from the same initial energy sources.*

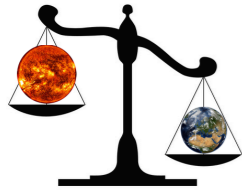
**The Earth atmosphere** [ Lovelock 1965; 1975. Hitchcock and Lovelock, 1967 ]

**Co-evolution of Earth geochemical cycles and life** [ Grenfell et al., 2010. Lammer et al., 2010 ]



*“Once candidate disequilibria are identified, alternative explanations must be eliminated. Life is the hypothesis of last resort” (Sagan et al., 1993)*





# Atmospheric Chemical Disequilibrium



## A PHYSICAL BASIS FOR LIFE DETECTION EXPERIMENTS

By DR. J. E. LOVELOCK

As yet, there is no direct evidence of life elsewhere in the solar system. However, an understanding of the physical and chemical environments of any of the planetary bodies...

It is not surprising in view of the rapid expansion of space probe experiments and of the considerable uncertainties already stated here, that the proposed experiments in life detection all ask the question: "Is there life as we know it?" Most certainly it is difficult to envisage in detail an alien biotechnology, it would seem pointless and very unscientific to send a space probe to detect a specific life form...

### Recognition of Life

It is a relatively simple matter to distinguish between living and inorganic matter on Earth by biochemical experiments even though no formal definition of life is biochemically tenable. Experiments suggest, for example, that a system capable of converting water...

On the basis of the physical phenomena already mentioned, a planet bearing life is distinguishable from a sterile one as follows: (1) The composition of inorganic molecules and of elements and of ions...

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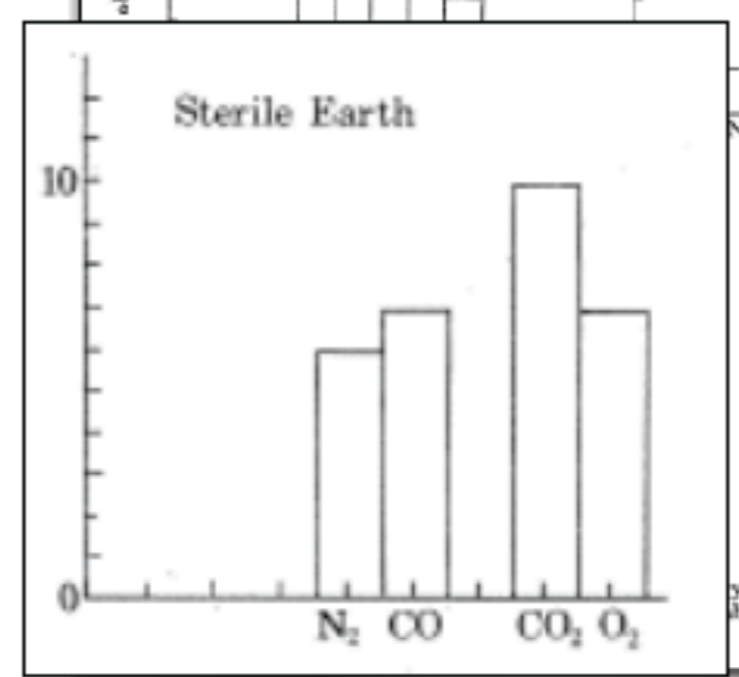
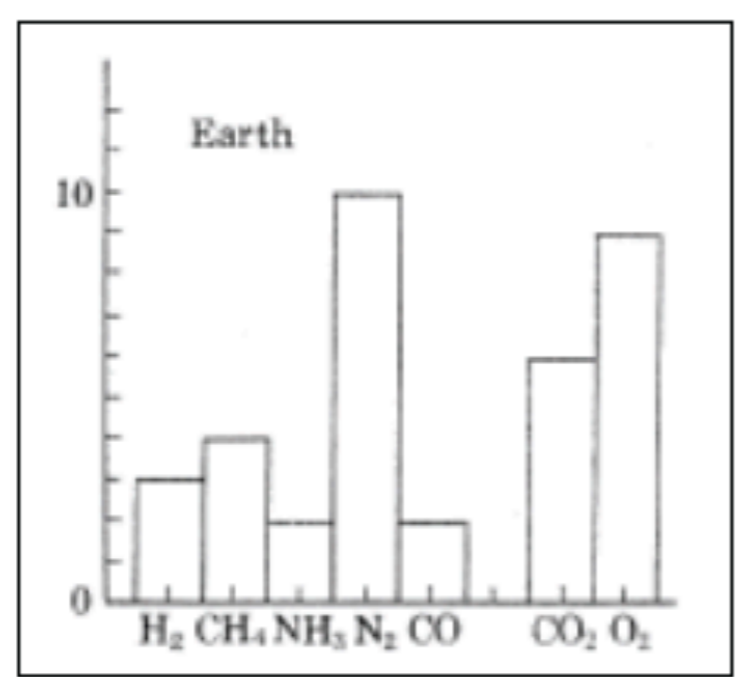
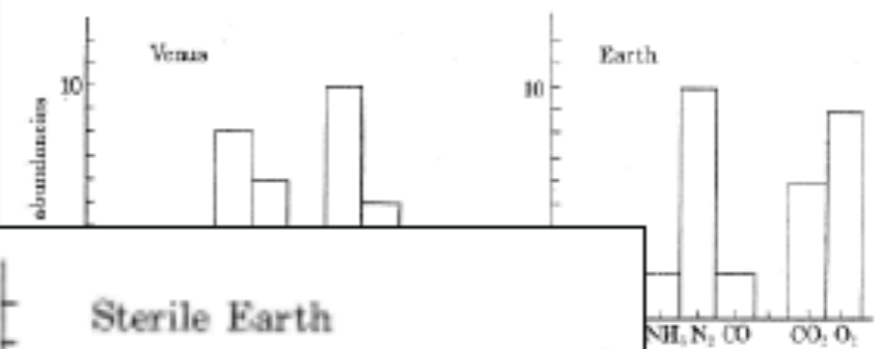
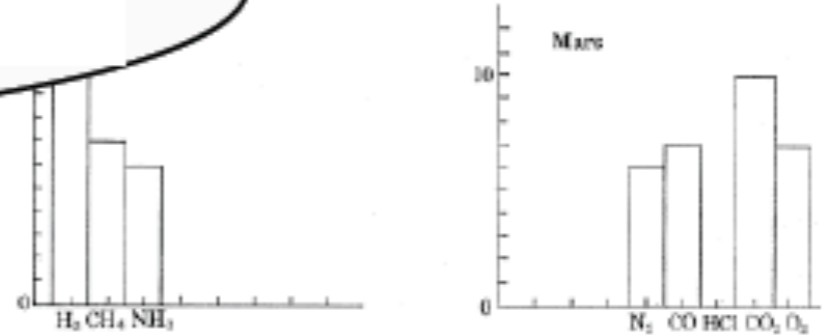
This relationship with chemical disequilibrium would be a characteristic but still recognizable signal to be expected to penetrate into the planetary surface and its past history as fossils and as rocks of biological origin.

### Experiments for Detection of Life

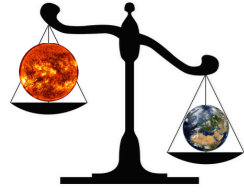
The distinguishing features of a life-bearing planet, described here, suggest the following simple experiments to determine if life:

(1) Search for order. (2) Order by chemical structure and sequence of structure. A simple gas chromatograph or a mass spectrometer instrument would, with careful molecular selection as well as detailed identification...

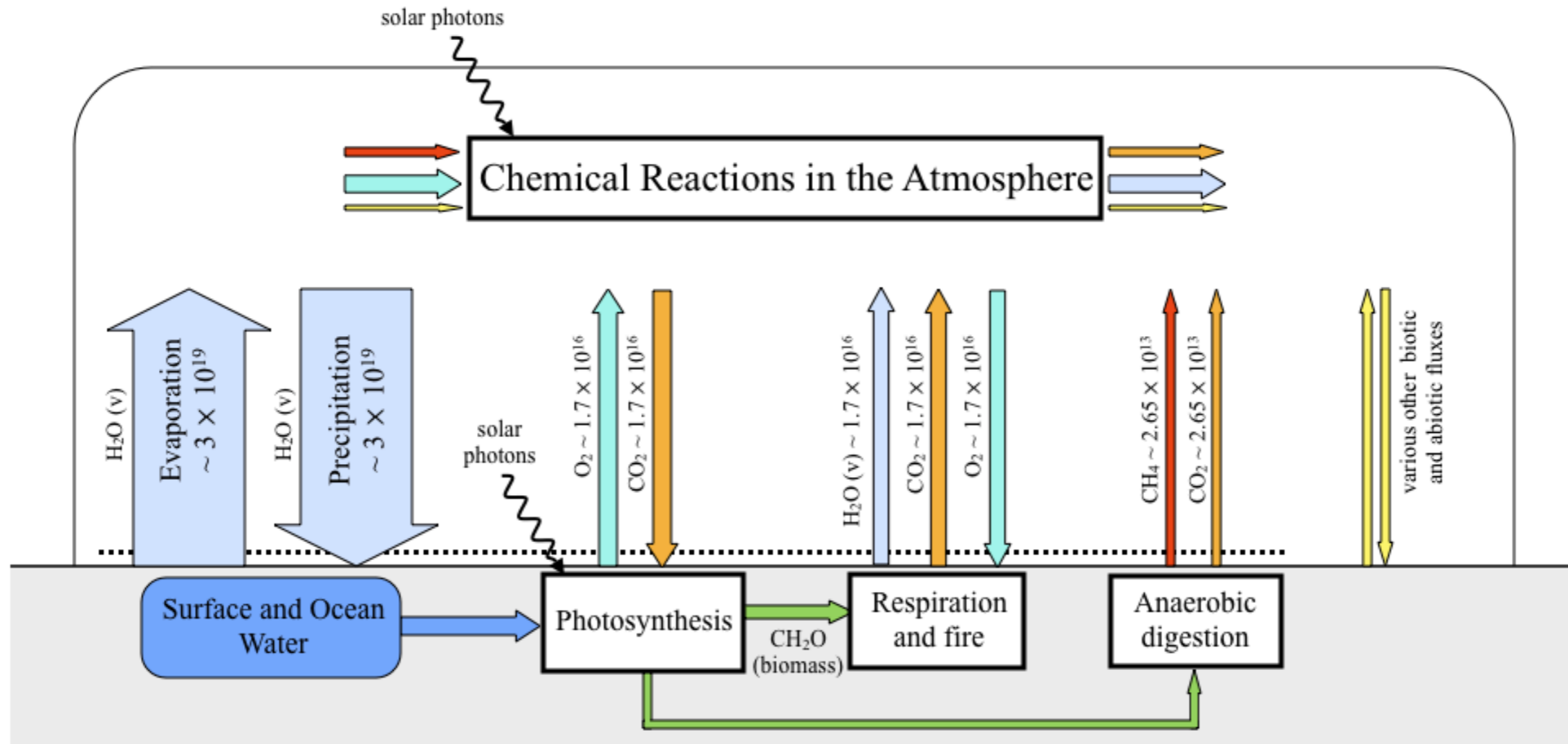
J. E. Lovelock (Discussion Meeting)



by atmospheres. abundances.



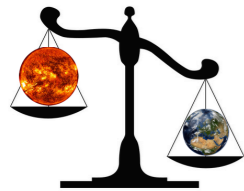
# Earth's methane disequilibrium



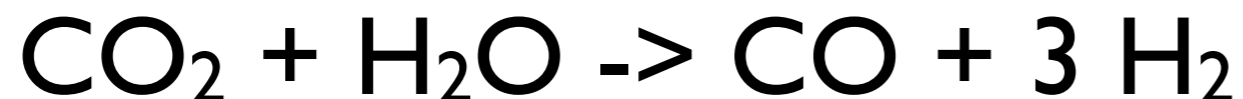
The contemporaneous presence of  $O_2$  and  $CH_4$  into the Earth's atmosphere is maintained by a power of  $\sim 0.67$  TW

About 0.43 TW are given by living processes (animal enteric fermentation, 0.13TW; rice paddies 0.09TW).





# Atmospheric Chemical Disequilibrium



$$\alpha(f_i, P) = \frac{f_{\text{CO}} f_{\text{H}_2\text{O}}}{f_{\text{CO}_2} f_{\text{H}_2}^3 P^2} = K_{eq}(T)$$

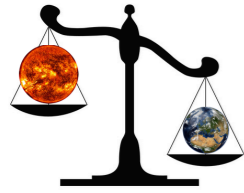
$$K_{eq}(T) = e^{-\Delta G/RT} = e^{-(\Delta H/RT - \Delta S/R)}$$

Line, M. R., Yung Y. L., ApJ 2013a

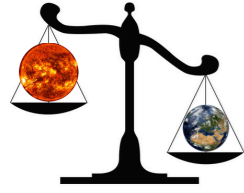
Line, M. R., Yung Y. L., ApJ 2013b

Line, M. R., Yung Y. L., ApJ 2013c

C/O ratio - Madhusudhan N., ApJ, 758:36, 2012



# How to calculate (and compare) disequilibrium in chemical processes

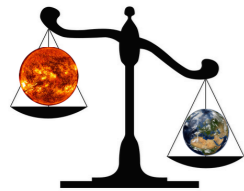


## *The extent of chemical disequilibrium*

In order to measure the extent of disequilibrium, we have to deal with the thermodynamics of non-equilibrium (irreversible) processes.

The distance of a system from its equilibrium condition (i.e. the measure of its irreversibility) is given by the entropy production within a system:

$$d_i S/dt$$



# The extent of chemical disequilibrium



$$\frac{d_i S}{dt} = J \cdot X = \frac{d\xi}{dt} \cdot \frac{\alpha}{T}$$

Extent of reaction:

$$\xi(t) = \frac{[A]_0 - [A](t)}{\nu_A}$$

Chemical Affinity

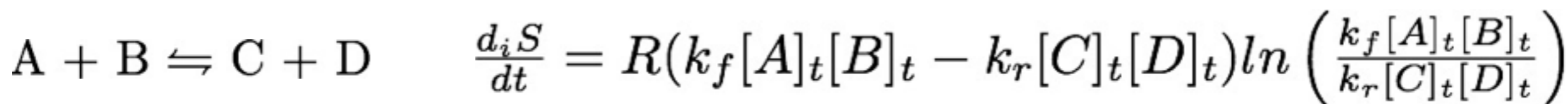
$$\alpha(t) = - \left( \frac{\partial \Delta_r G(t)}{\partial \xi} \right)_{T,p}$$

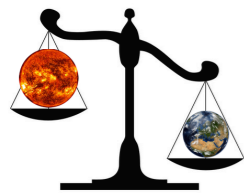
It can be also written as:

$$\frac{d_i S}{dt} = R \cdot (R_f - R_r) \cdot \ln \left( \frac{R_f}{R_r} \right)$$

$R_f$  = forward rate

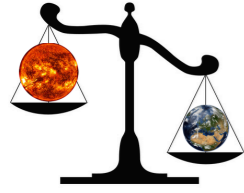
$R_r$  = backward rate





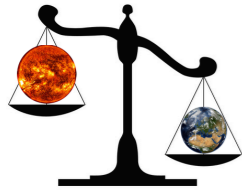
- Python pre-processor provides Fortran routines
- Creates modules from chemical network
- Dust evolution, cooling heating photoionization
- Large test suite
- Highly optimized, fast solvers
- Open source, bitbucket community
- Grassi T. et al., MNRAS 2014. doi:10.1093/mnras/stu114

[www.kromepackage.org](http://www.kromepackage.org)



# Atmospheric disequilibrium of the Earth

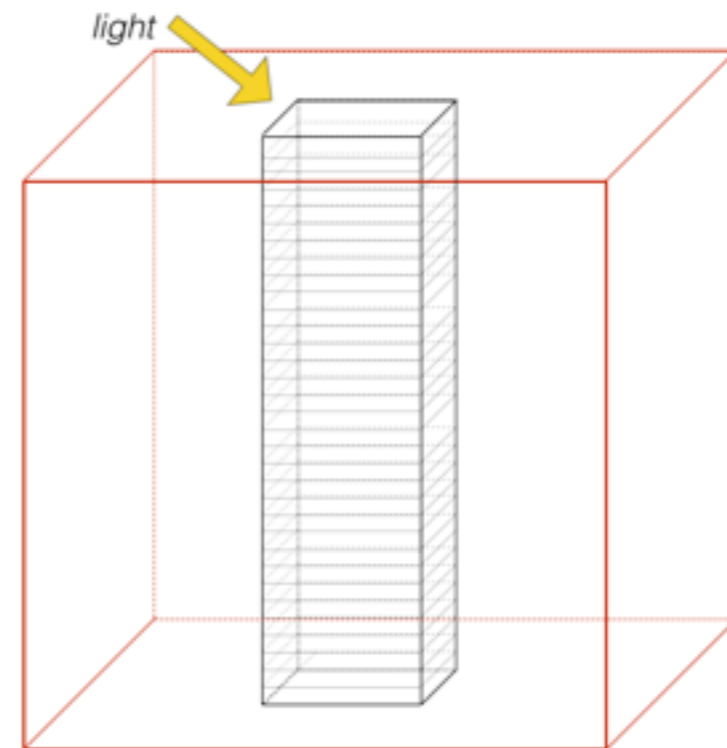
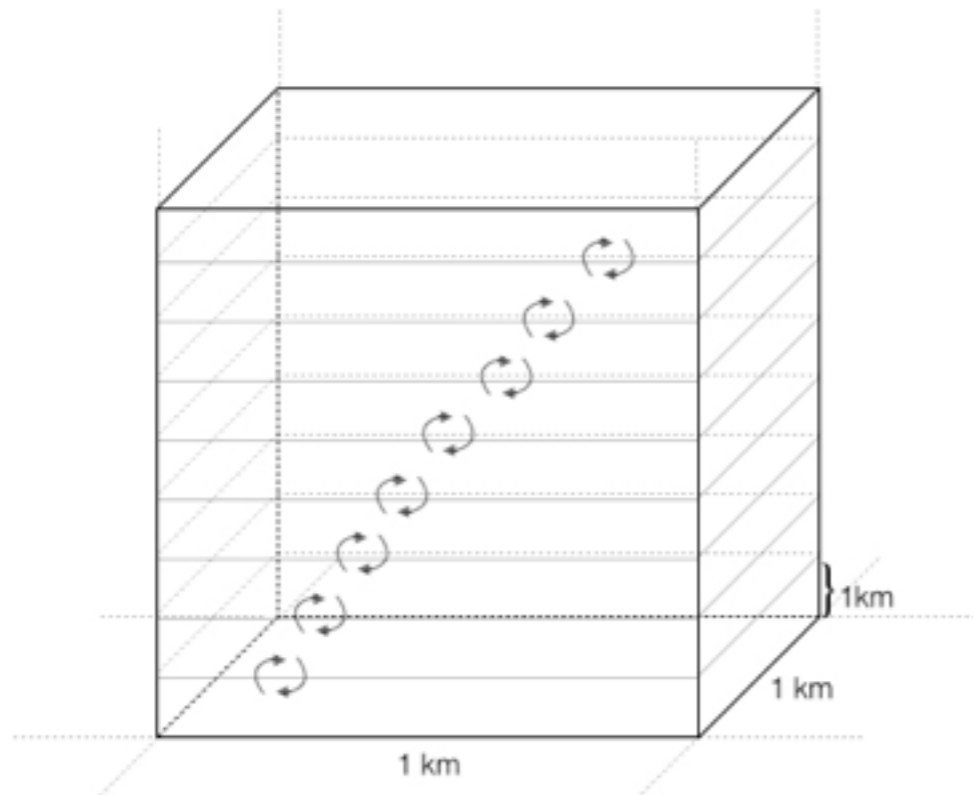




# Earth Atmospheric Chemical Disequilibrium



- \* Model: Kasting, J. F., and Donahue, T. M., J. Geophys. Res., 85,3255-3263. 1980 (K-80);
- \* 64 layers (~1km each);
- \* Eddy diffusion;
- \* Entropy production and the power dissipation:



$$\sigma = \frac{d_i S}{dt}$$

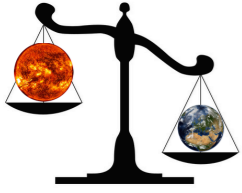
$$\frac{\sigma \times T}{A_{Earth}} \sim W m^{-2}$$

Simoncini, Brucato, Grassi, sub. to OLEB

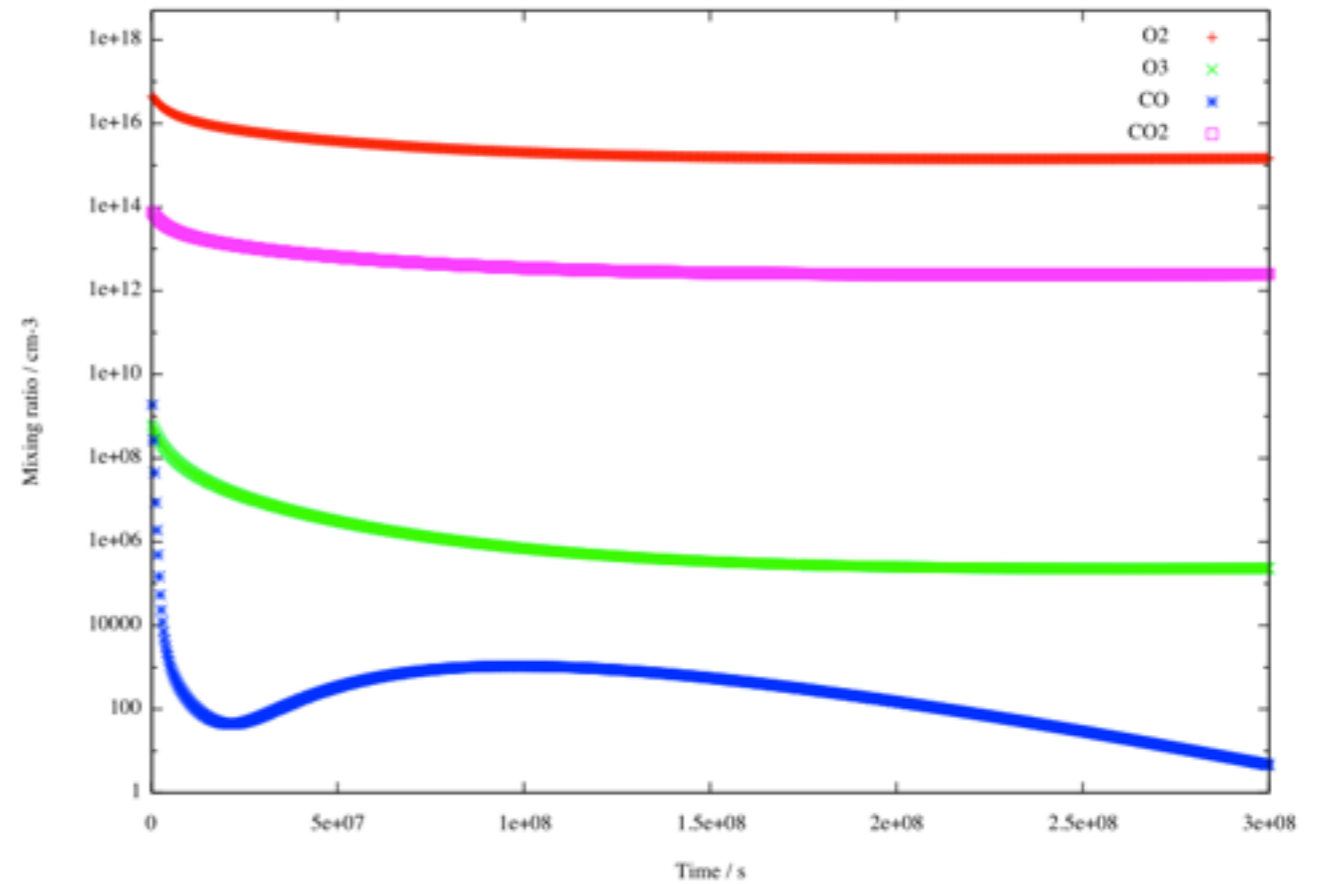
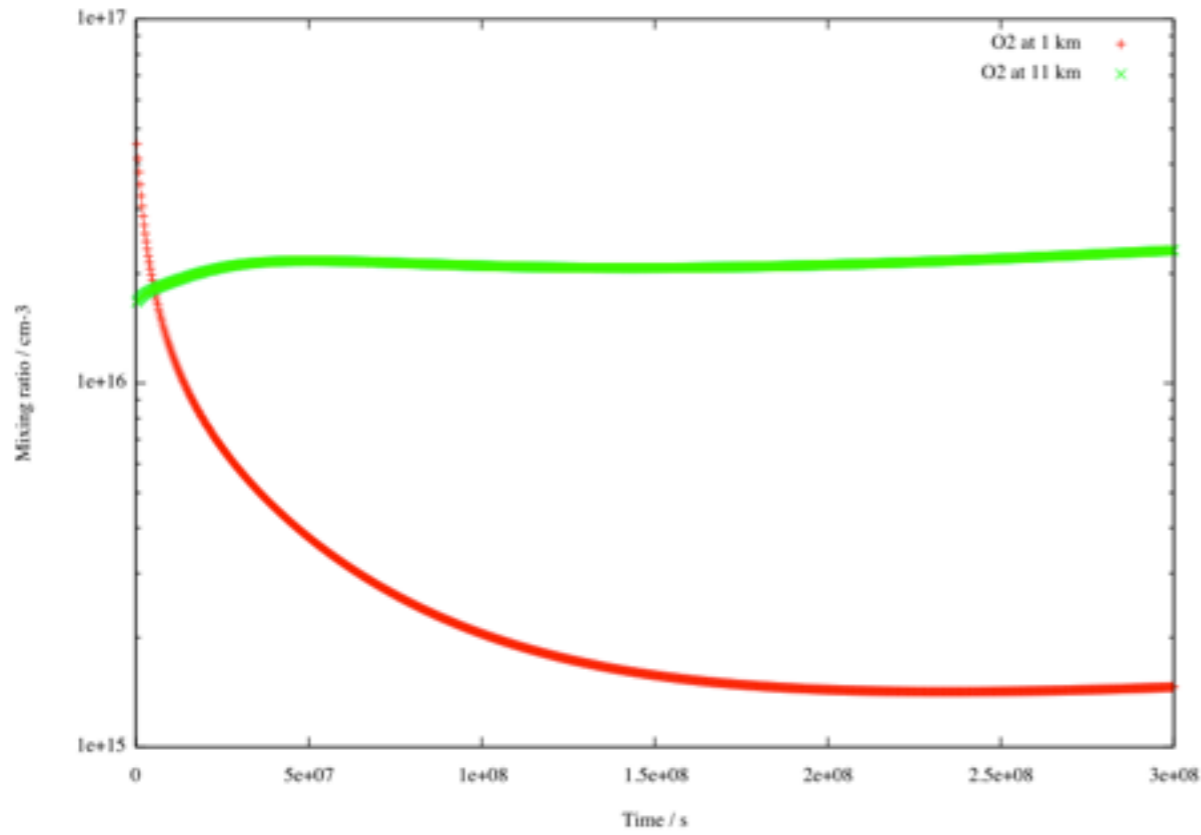
S. O. Danielache, E. Simoncini, Y. Ueno, Archean Atmospheres Modeled with the KROME Chemistry Package, JPGU 2014

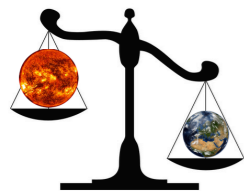
Simoncini E., Virgo N., Kleidon A., Quantifying drivers of chemical disequilibrium: theory and application to methane in the Earth's atmosphere. Earth System Dynamics 4, 1-15, 2013.

Angerhausen D., Sapers H., Simoncini E., and coworkers, An astrobiological experiment to explore the habitability of tidally locked M-Dwarf planets, IAU 2013 Proceedings.

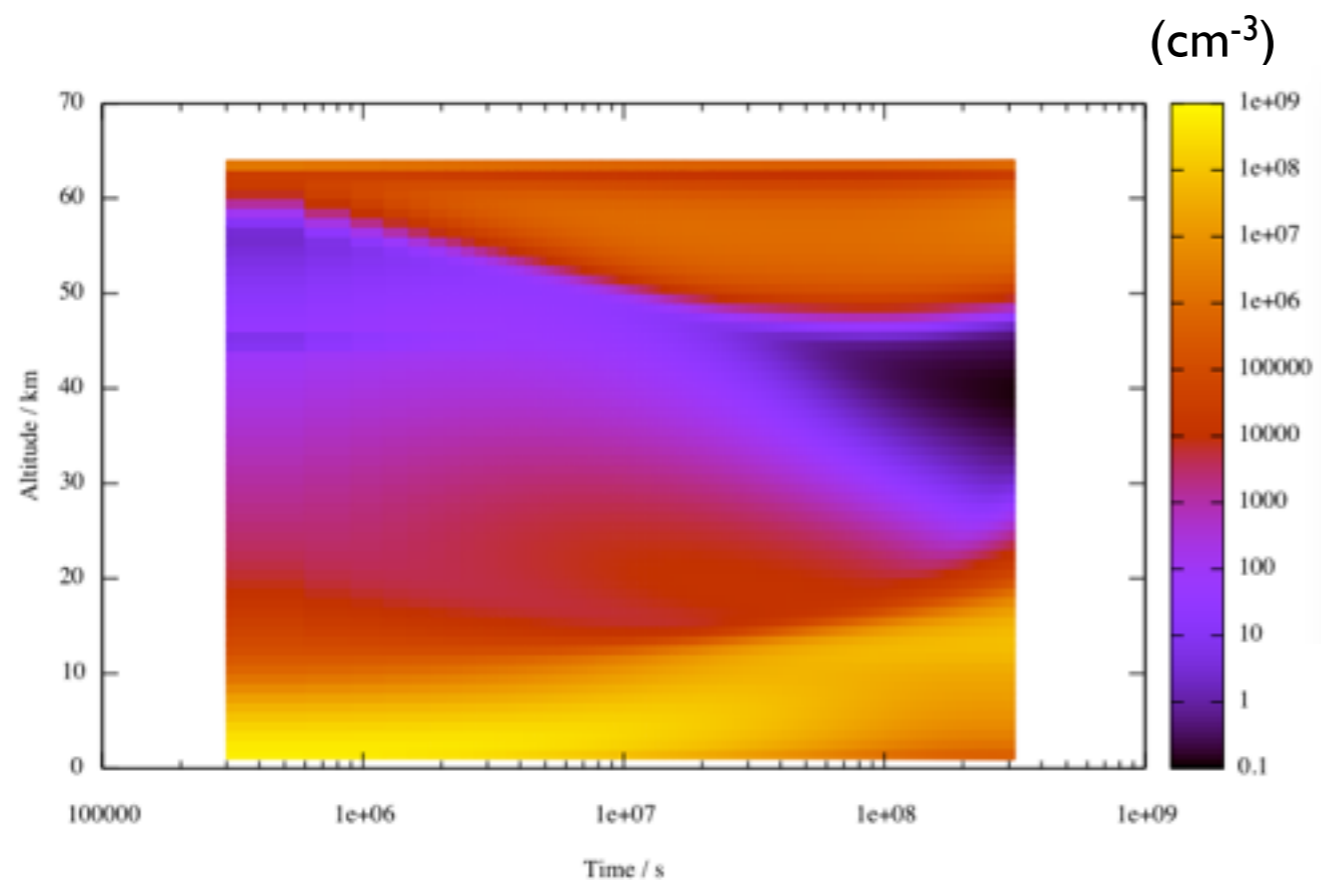


# Earth Atmospheric Chemical Disequilibrium

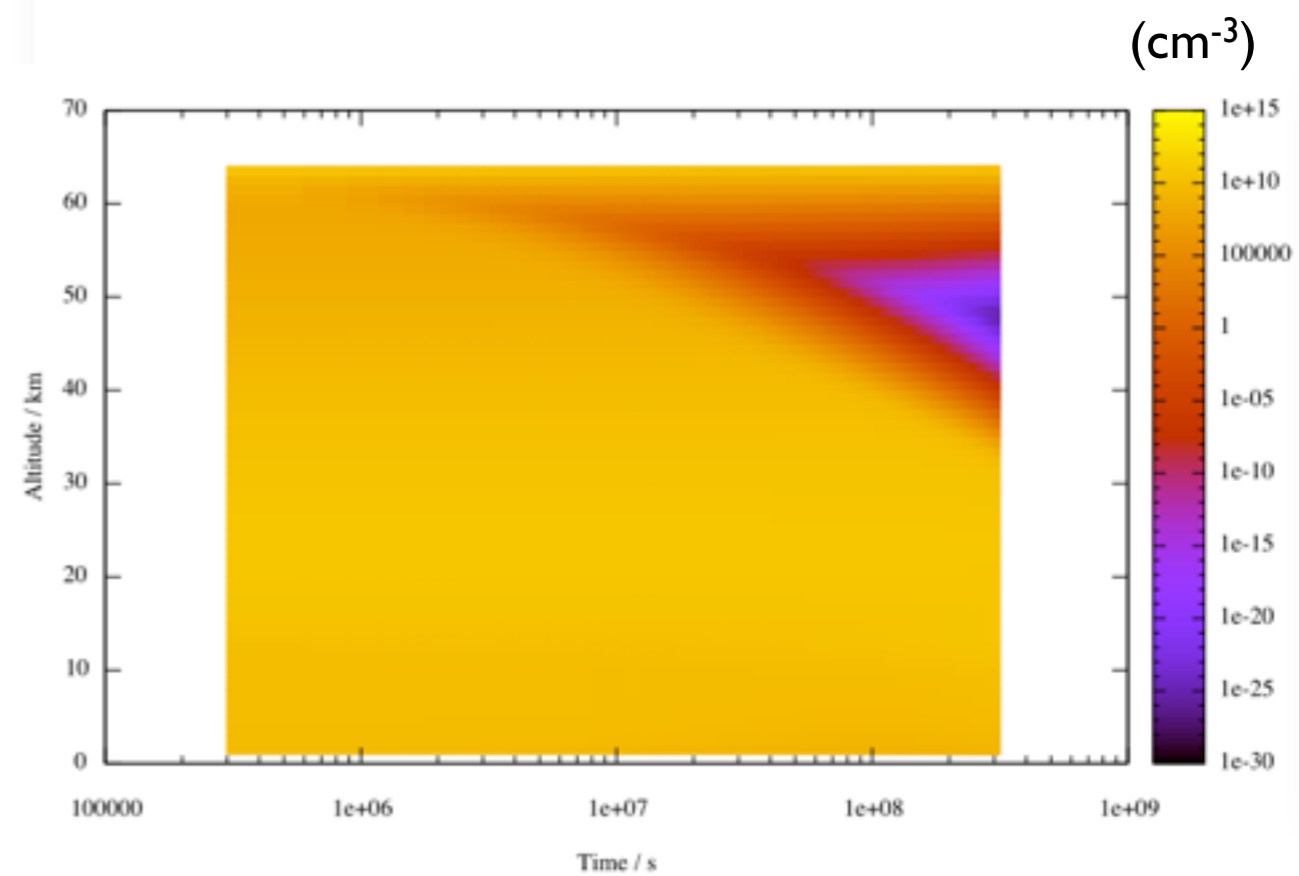


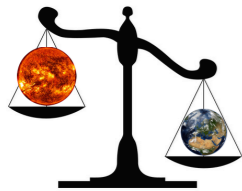


# Earth Atmospheric Chemical Disequilibrium



$\text{O}_3$

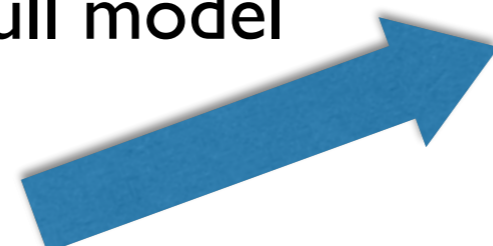




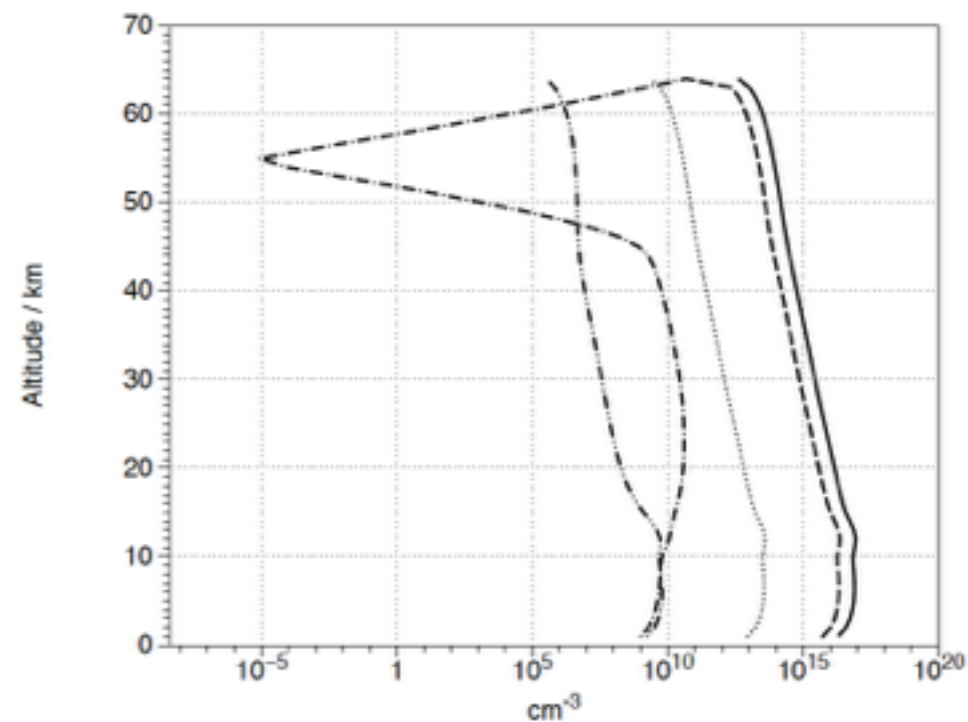
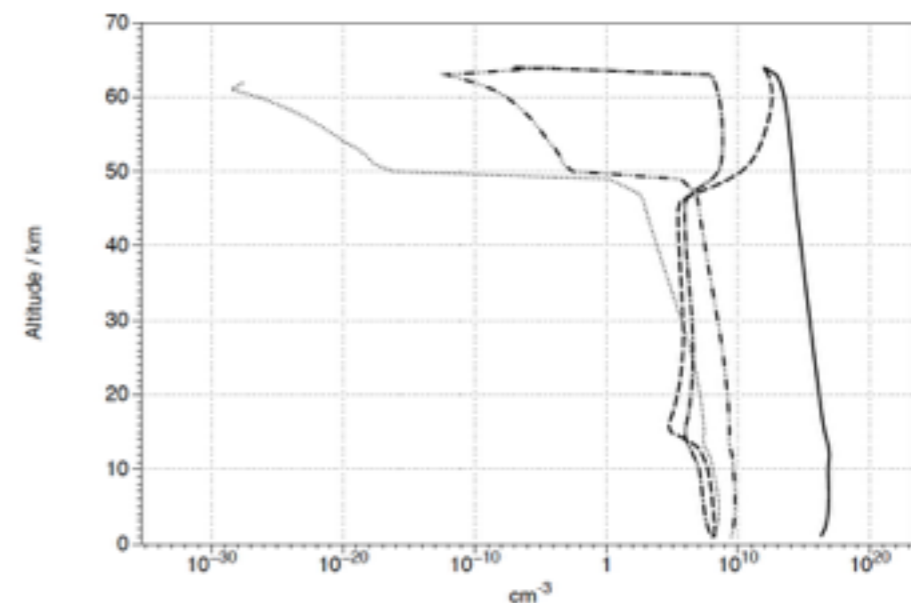
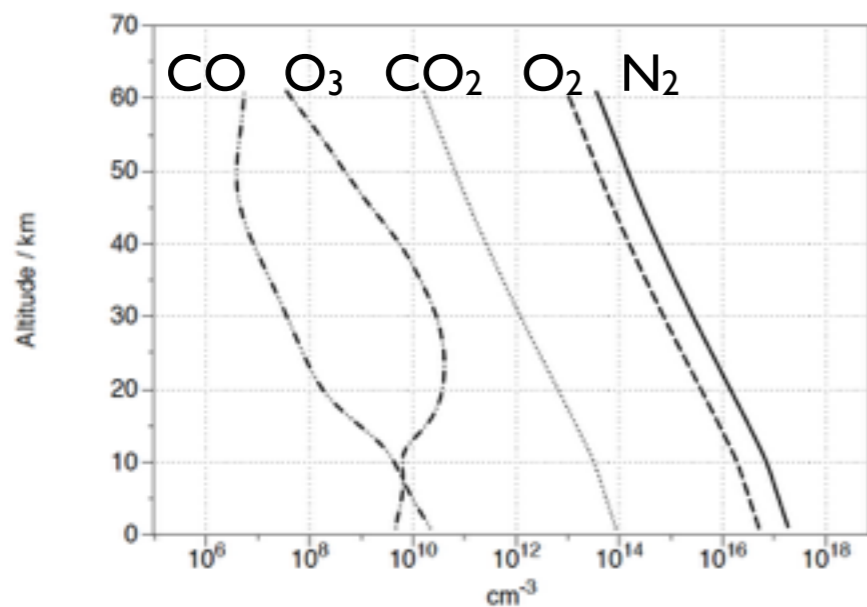
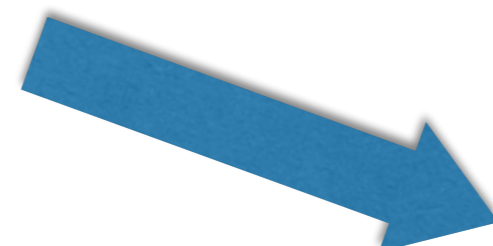
# Earth Atmospheric Chemical Disequilibrium

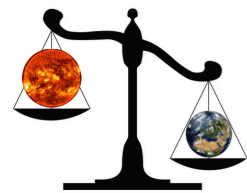


full model

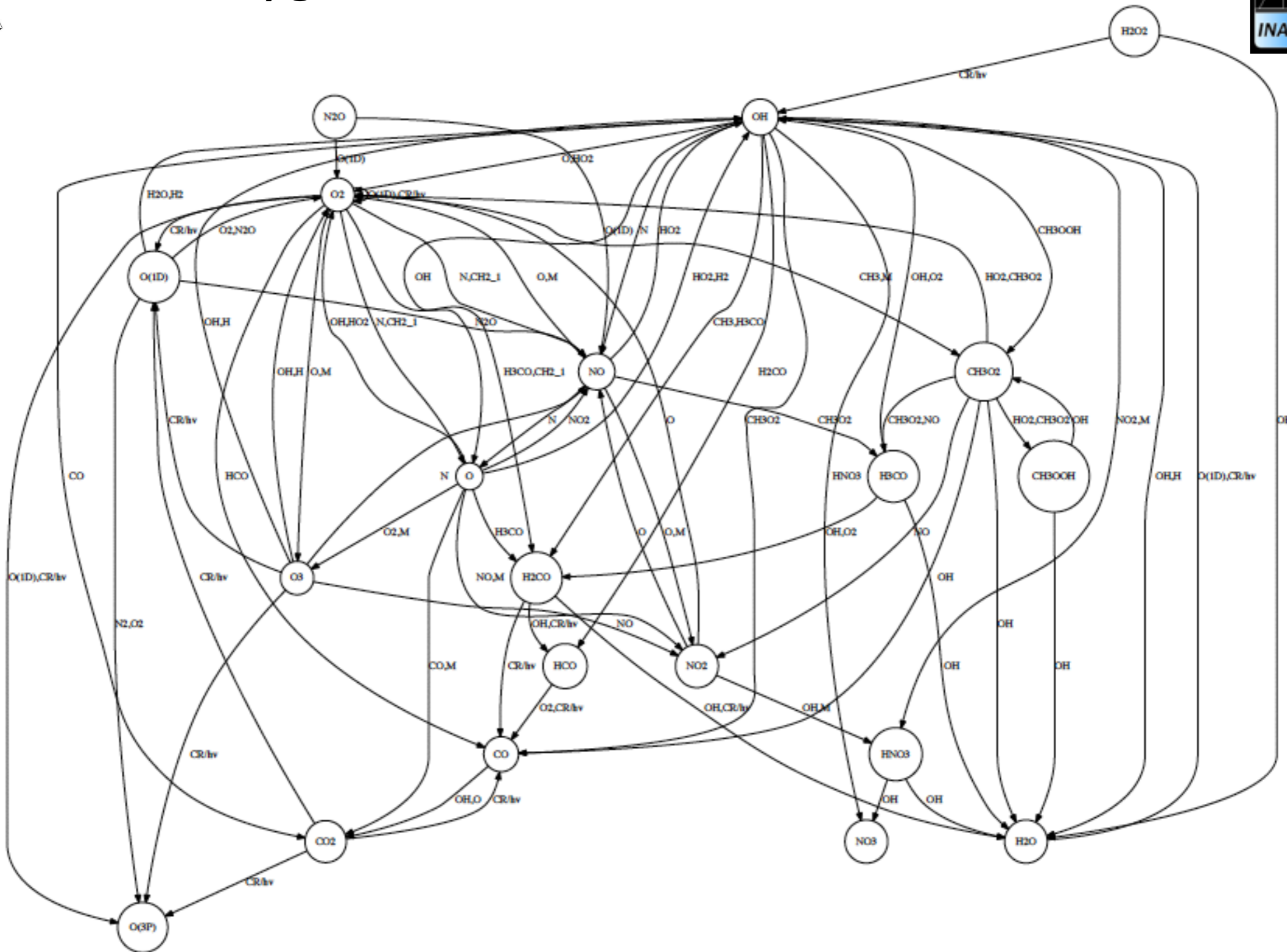


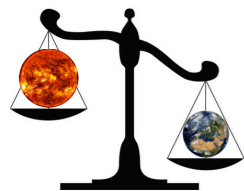
no  
photochemistry





# Oxygen reactions in the K-80 network



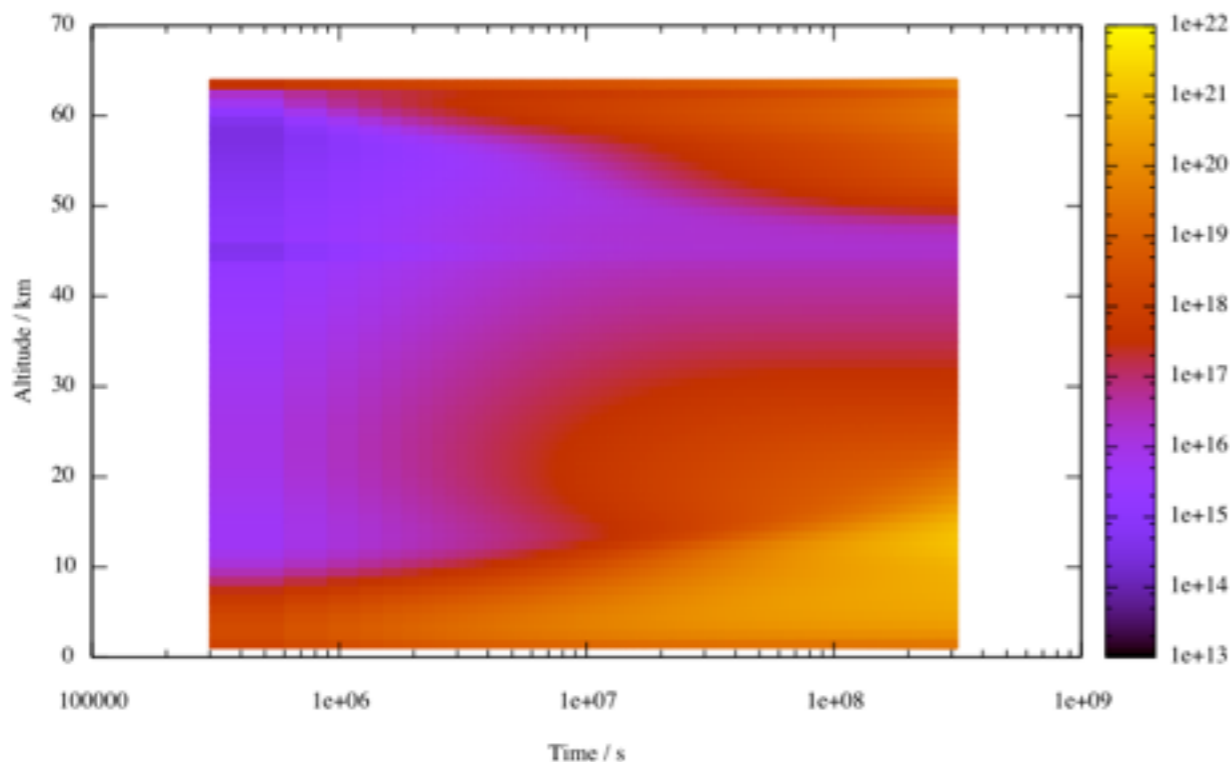


# Earth Atmospheric Chemical Disequilibrium

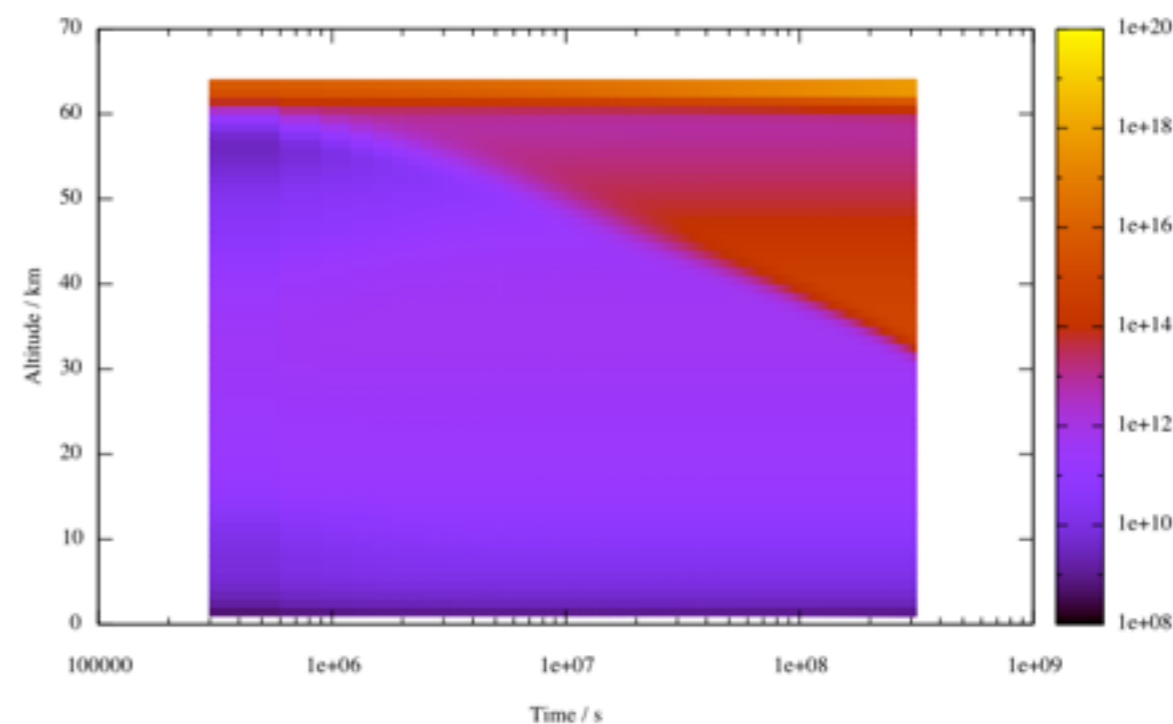


## *The energetic structure*

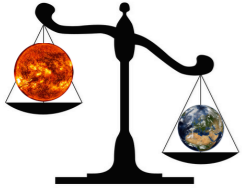
Present Earth (W)



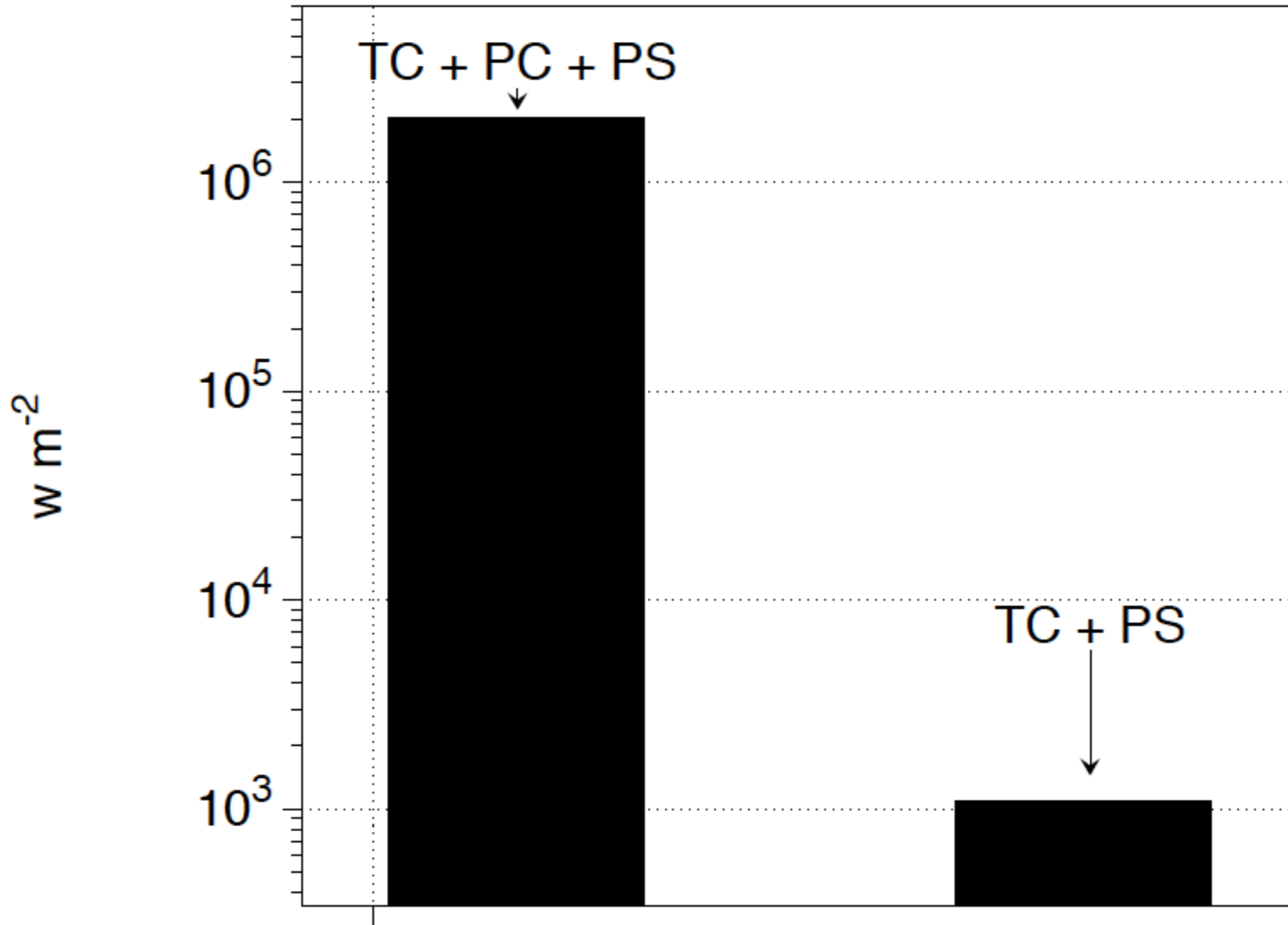
Present Earth without photochemistry (W)

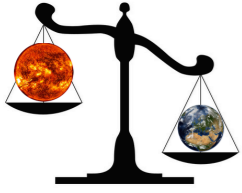






# Earth Atmospheric Chemical Disequilibrium

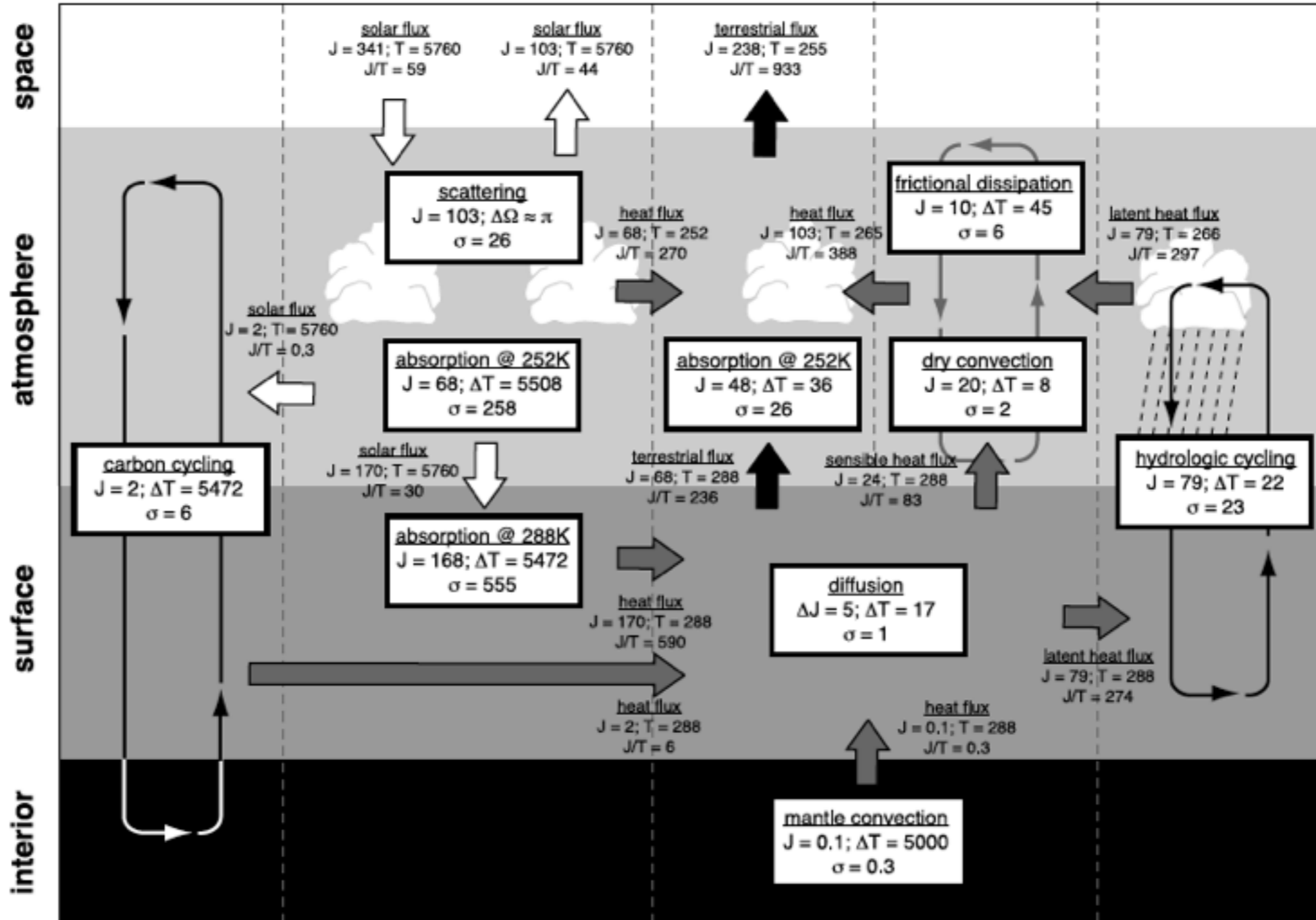




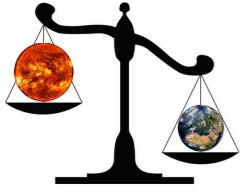
# Earth Powers



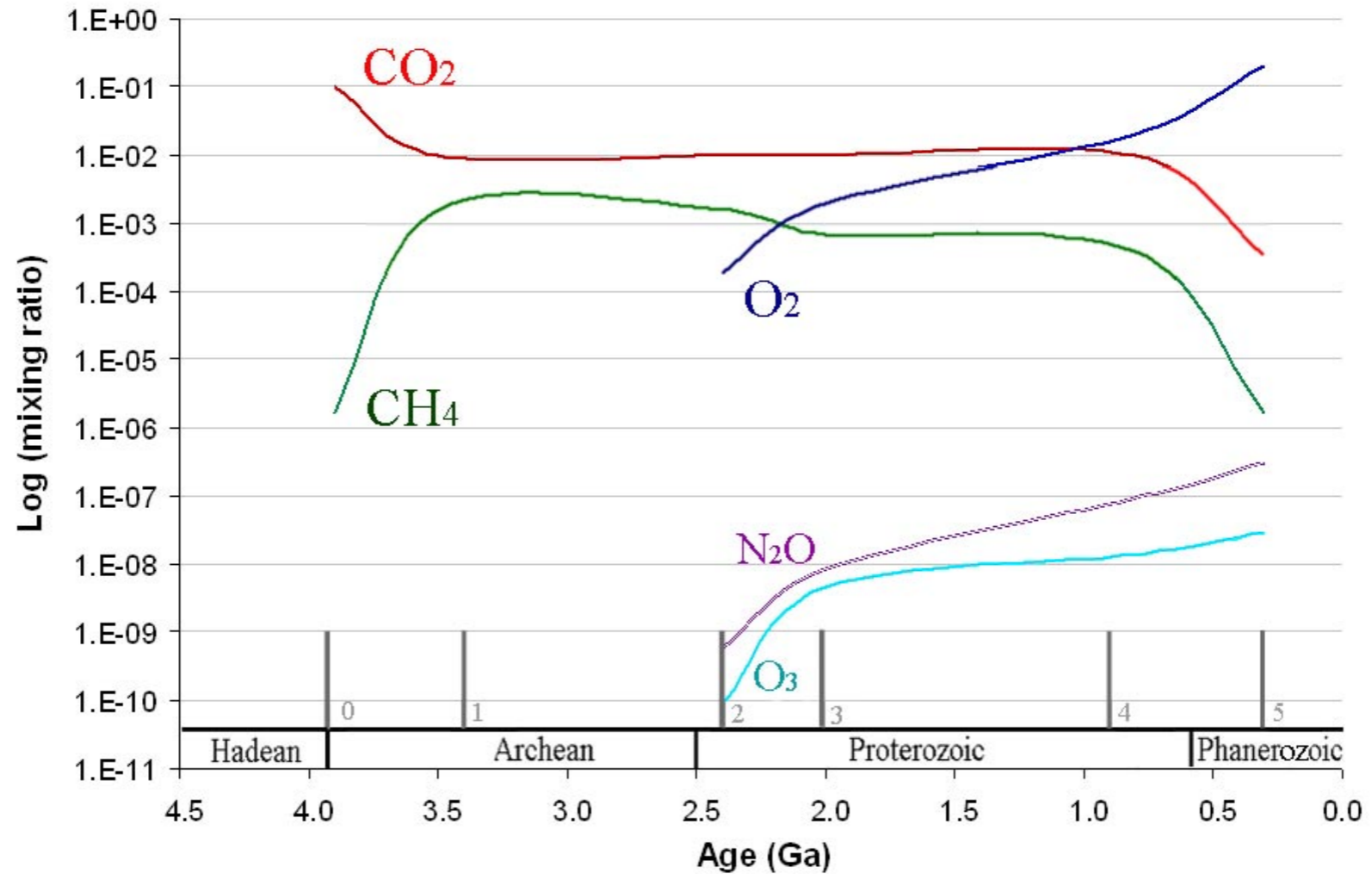
carbon cycle      solar radiation      terrestrial radiation      atmospheric dynamics      hydrologic cycle



$J \sim W m^{-2}$

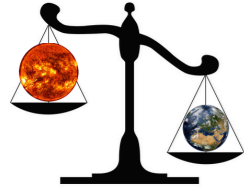


# Earth Atmospheric Chemical Disequilibrium



Epoch	Age (Ga)	CO <sub>2</sub> (mixing ratio)	CH <sub>4</sub> (mixing ratio)	O <sub>2</sub> (mixing ratio)	O <sub>3</sub> (mixing ratio)	N <sub>2</sub> O (mixing ratio)
0	3.9	1.00E-01	1.65E-06	0	0	0
1	3.5	1.00E-02	1.65E-03	0	0	0
2	2.4	1.00E-02	7.07E-03	2.10E-04	8.47E-11	5.71E-10
3	2.0	1.00E-02	1.65E-03	2.10E-03	4.24E-09	8.37E-09
4	0.8	1.00E-02	4.15E-04	2.10E-02	1.36E-08	9.15E-08
5	0.3	3.65E-04	1.65E-06	2.10E-01	3.00E-08	3.00E-07

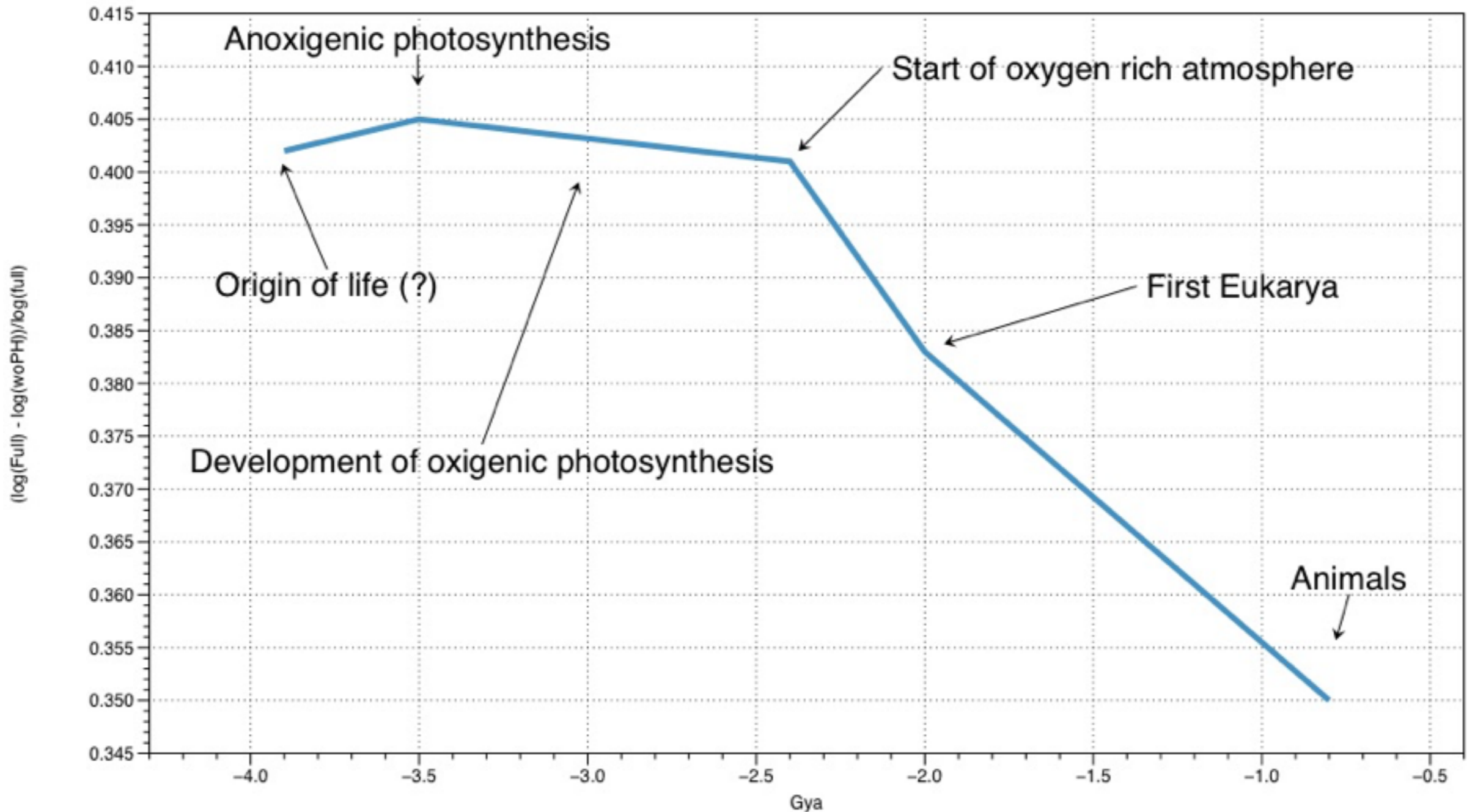
Simoncini, Brucato, Grassi, sub. to OLEB  
 Kaltenecker et al., *Apj* 658, 598, 2007  
 Kasting, J. F., *Scientific American Magazine*; 80 2004

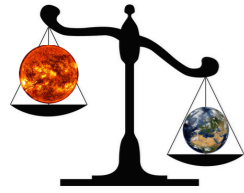


# Earth Atmospheric Chemical Disequilibrium



## Life origin and development *The weight of photochemistry*





# *Atmospheric extent of disequilibrium*



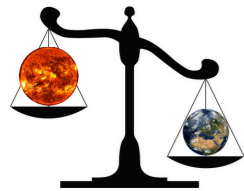
## Further studies

- > Earth + fluxes (steady state)
- > Earth + simplified biosphere (not stable LV model)
- > Analysis of reaction pathways
- > Deeper analysis of sulfur chemistry
- > Influence of Sun luminosity variability



*basis for habitability studies*

- > Atmospheric spectra
- > Mars atmosphere
- > Rocky and warm/hot exoplanets  
(new models)
- > Other Solar System planets and moons



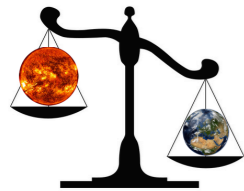
## *KROME + Planetary Atmosphere Applications*

# Modeling exoplanets with KROME

+ *Y. Miguel*

*Observatoire de Côte d'Azur, Nice, France*





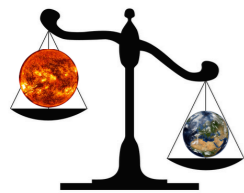
# Hot rocky exoplanet



- \*  $M = 10 \cdot M_e$
- \* Star = Sun ( $T_{\text{superf}} = 5777 \text{ K}$ )
- \*  $T \sim 2200 \text{ K}$  (temperature profile by Ito et al., *ApJ* **801** 144, 2015)
- \* 30 species from accretion (Y. Miguel code)  
—> part of them condensates
- \* Build-up abundances and pressure profile using the Scale Height:

$$H = \frac{k_b T}{Mg}$$

- \* Using the  $T, p$  conditions, take out not gaseous species
- \*  $K_{zz} = 10^6$  (low Martian atmosphere)
- \* Build-up the network: NIST, KIDA, exoplanets literature, Mercury atmosphere



# Hot rocky exoplanet: effect of diffusion

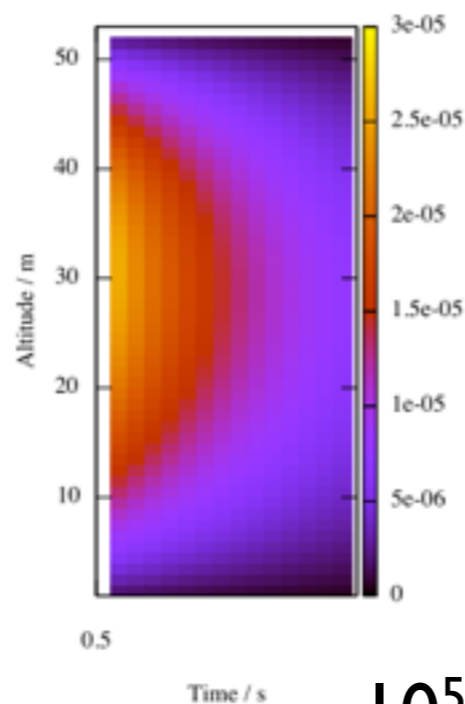
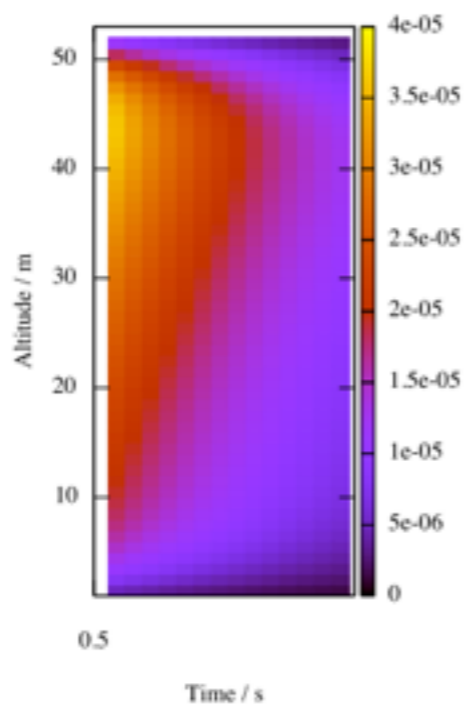
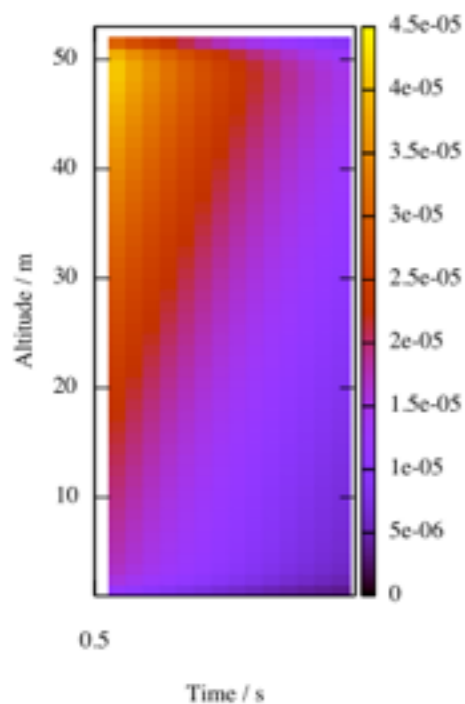


$10^5$

$10^6$

$10^7$

$$K_{zz} \sim \text{cm}^2 / \text{s}$$

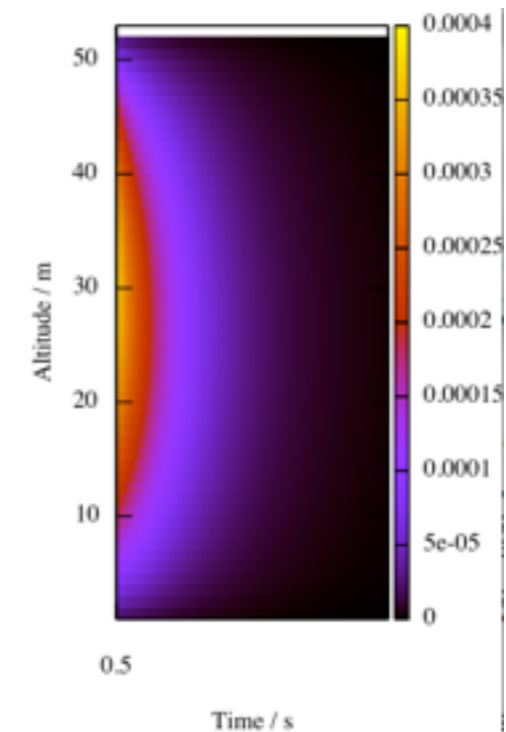
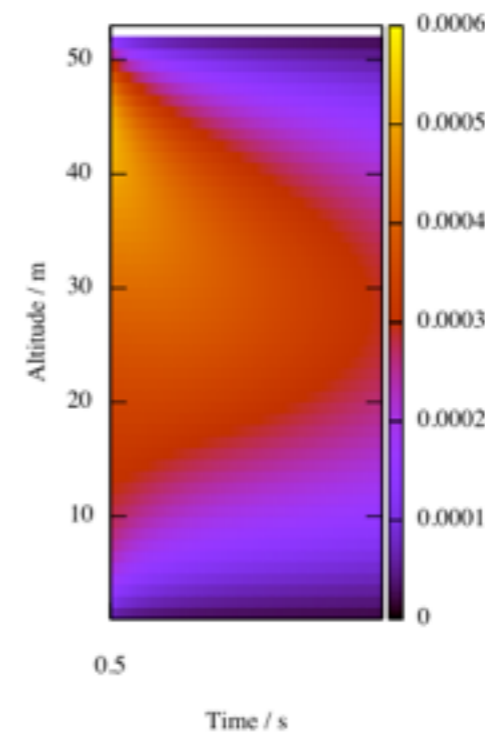
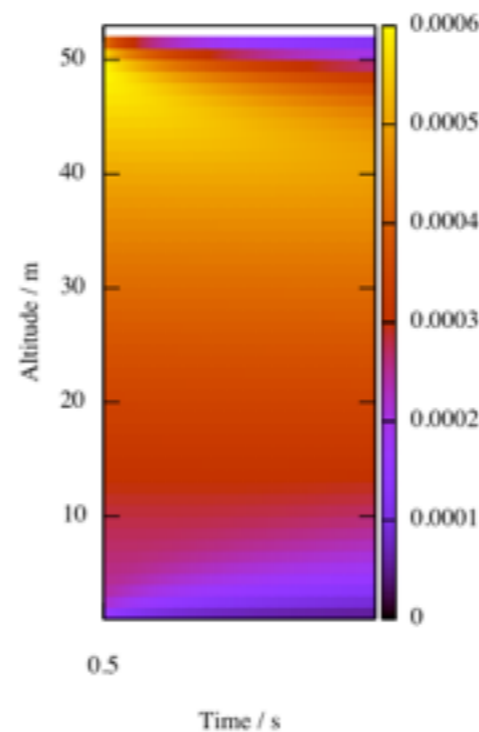


O<sub>2</sub>

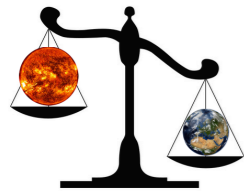
$10^5$

$10^6$

$10^7$



Na



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J. R. Brucato, Astrophysical Observatory of Arcetri - INAF, Firenze, Italy

T. Grassi, Starplan/NBI, University of Copenhagen, Denmark

S. O. Danielache, Sophia University, Tokyo, Japan

Y. Miguel, Observatoire Côte d'Azur, Nice, France

A. Chiavassa, Observatoire Côte d'Azur, Nice, France

M. J. Russell, JPL, CalTech-NASA, Pasadena, CA, USA

A. Kleidon, Max Planck Institute for Biogeochemistry, Jena, Germany

*...And all the members of the TDE Focus Group*